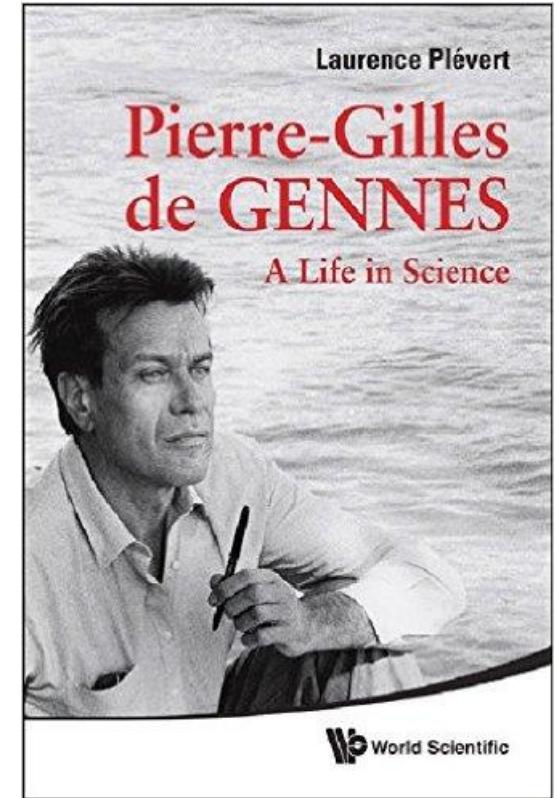


Part III

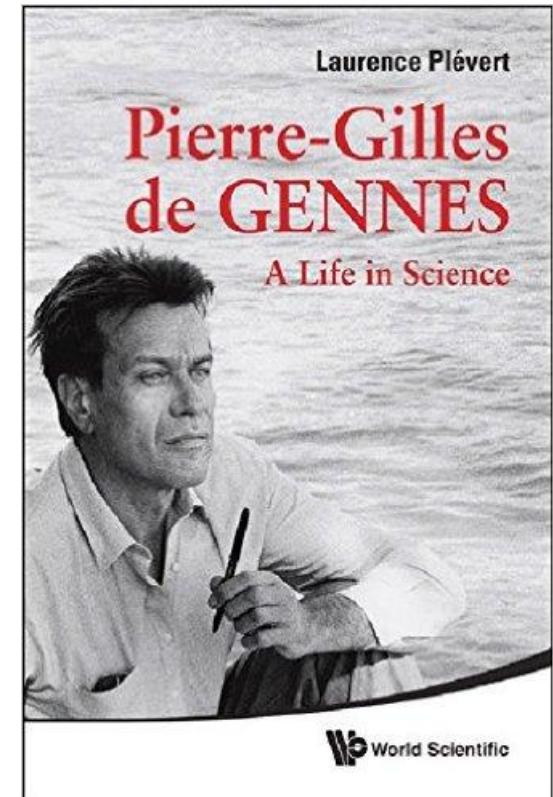
Пьер Жиль де Жен (фр. *Pierre-Gilles de Gennes*; 24 октября 1932, Париж — 18 мая 2007, Орсе) — французский физик, лауреат Нобелевской премии по физике в 1991 году «за обнаружение того, что методы, развитые для изучения явлений упорядоченности в простых системах, могут быть обобщены на жидкие кристаллы и полимеры». Де Жен известен прежде всего тем, что открыл структуру, положившую начало производству ЖК-дисплеев. За множество фундаментальных открытий многие научные круги называют де Жена «Ньютоном нашего времени».



Пьер Жиль де Жен (фр. *Pierre-Gilles de Gennes*: 24 октября 1932, Париж —

**“The easiest thing
in physics is the
mathematics, the
difficult bit is what
it means”**

множество фундаментальных открытий многие научные круги называют де Жена «Ньютоном нашего времени».



дисплеев. За

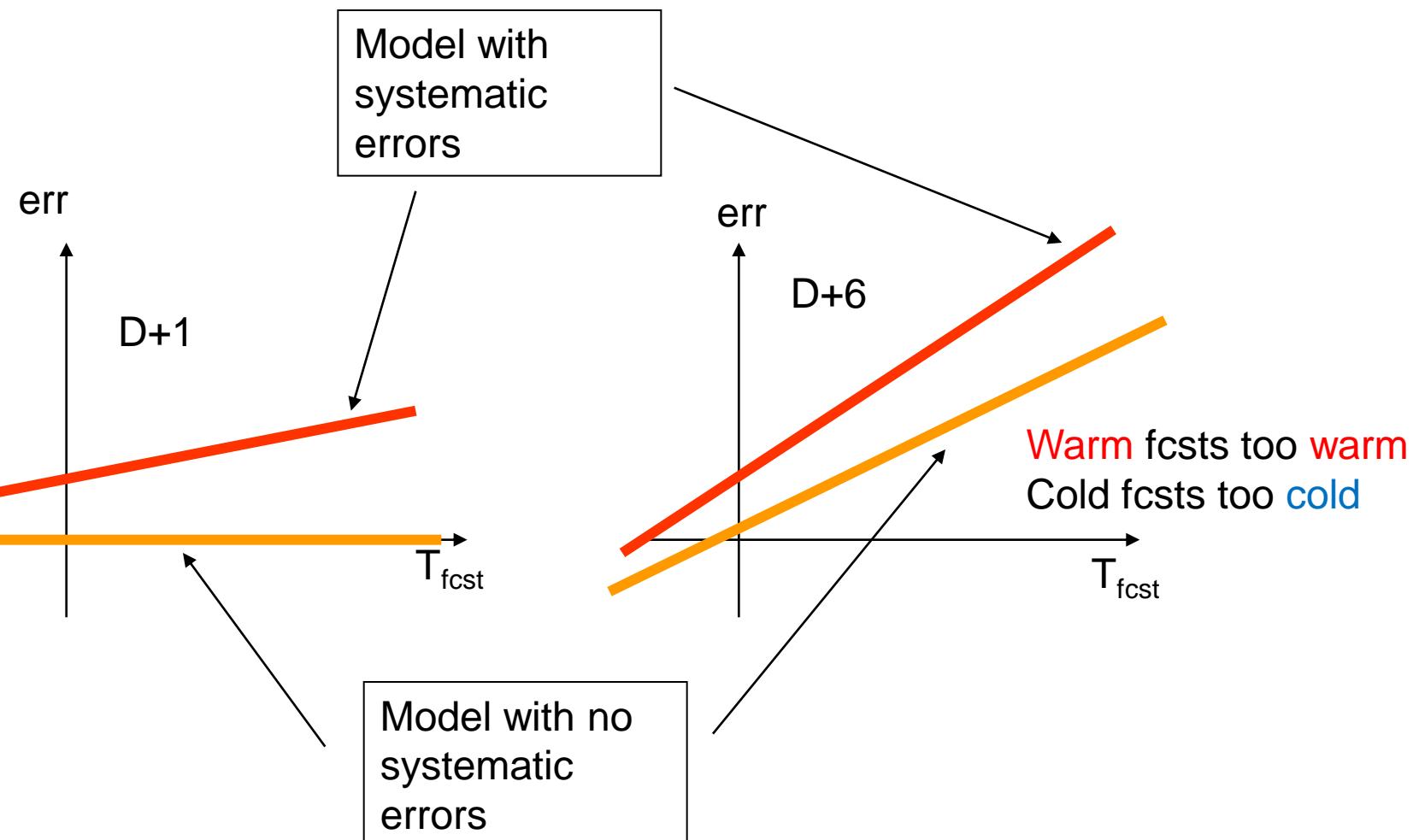
Mean errors are not the only “systematic” errors

- a) A mean error which is much smaller than the variance does not point to any “systematic” error
- b) In the opposite case we are dealing with a plain bias. But that is not the only type of “systematic” error
- c) Errors are systematic if they can be described/explained by a (linear) equation

False systematic errors

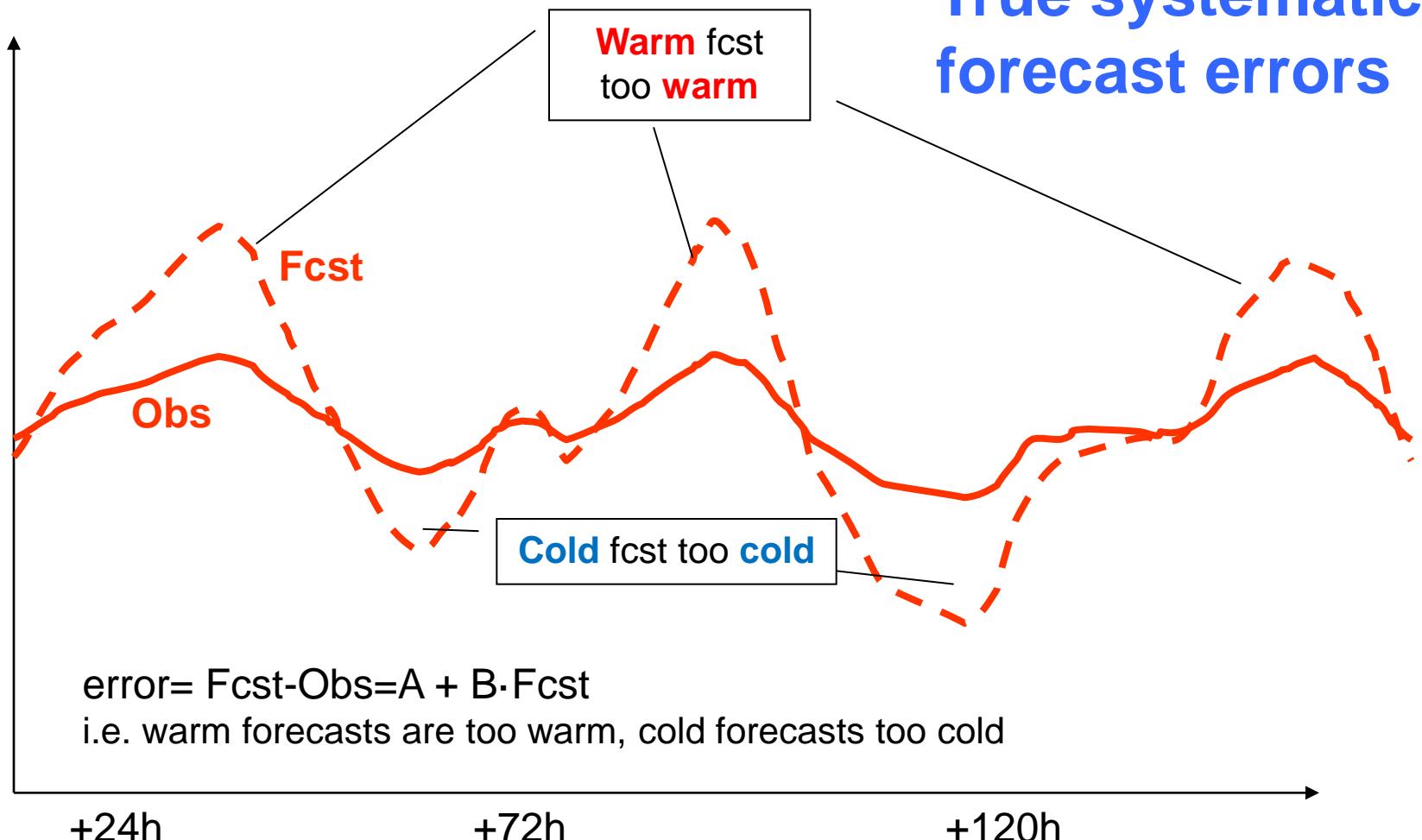
Non-systematic errors that look like
systematic

A model with no systematic errors (orange) will, with increasing lead time, appear to show a false drift towards increasing systematic errors

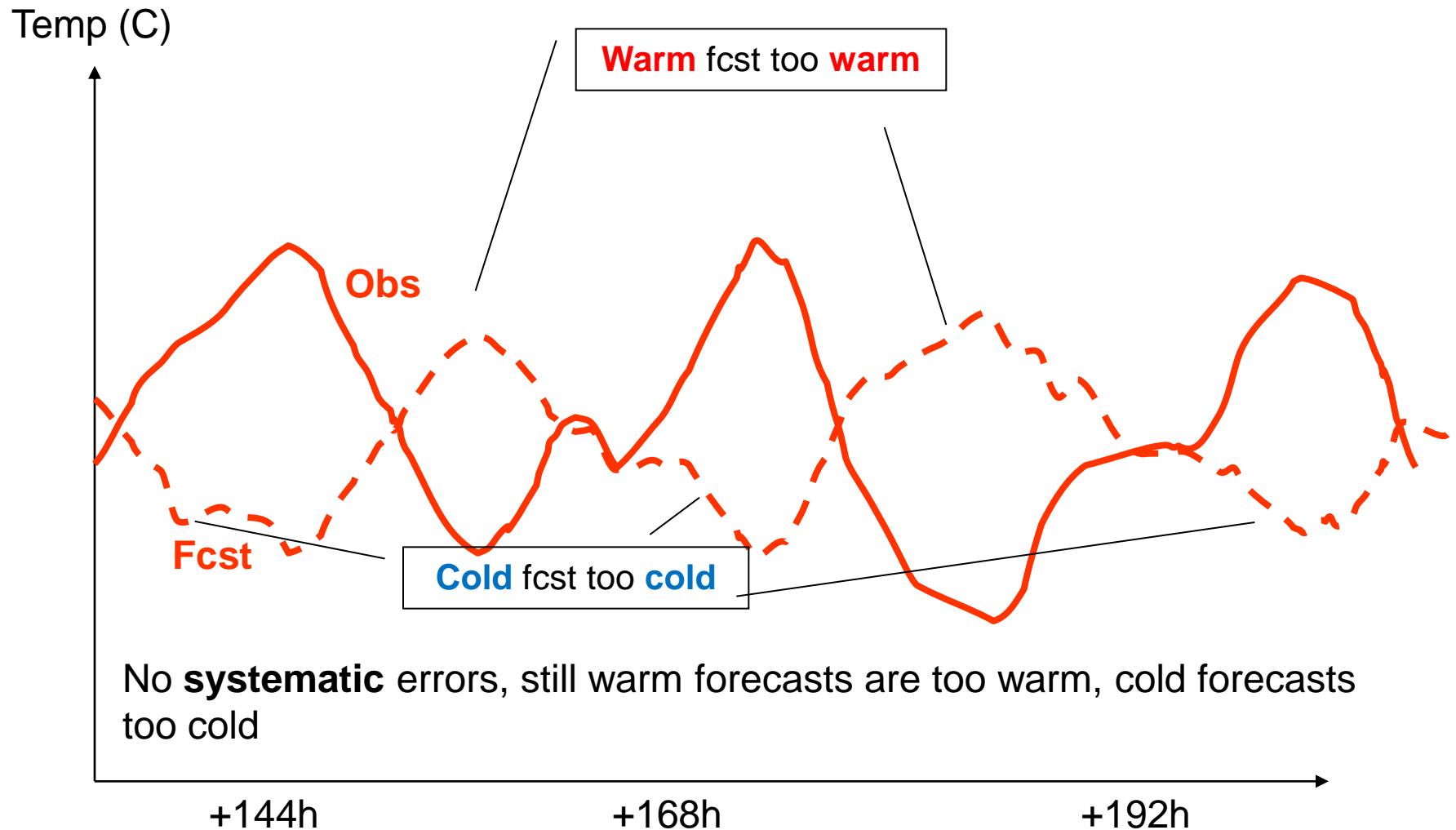


Temp (C)

True systematic forecast errors



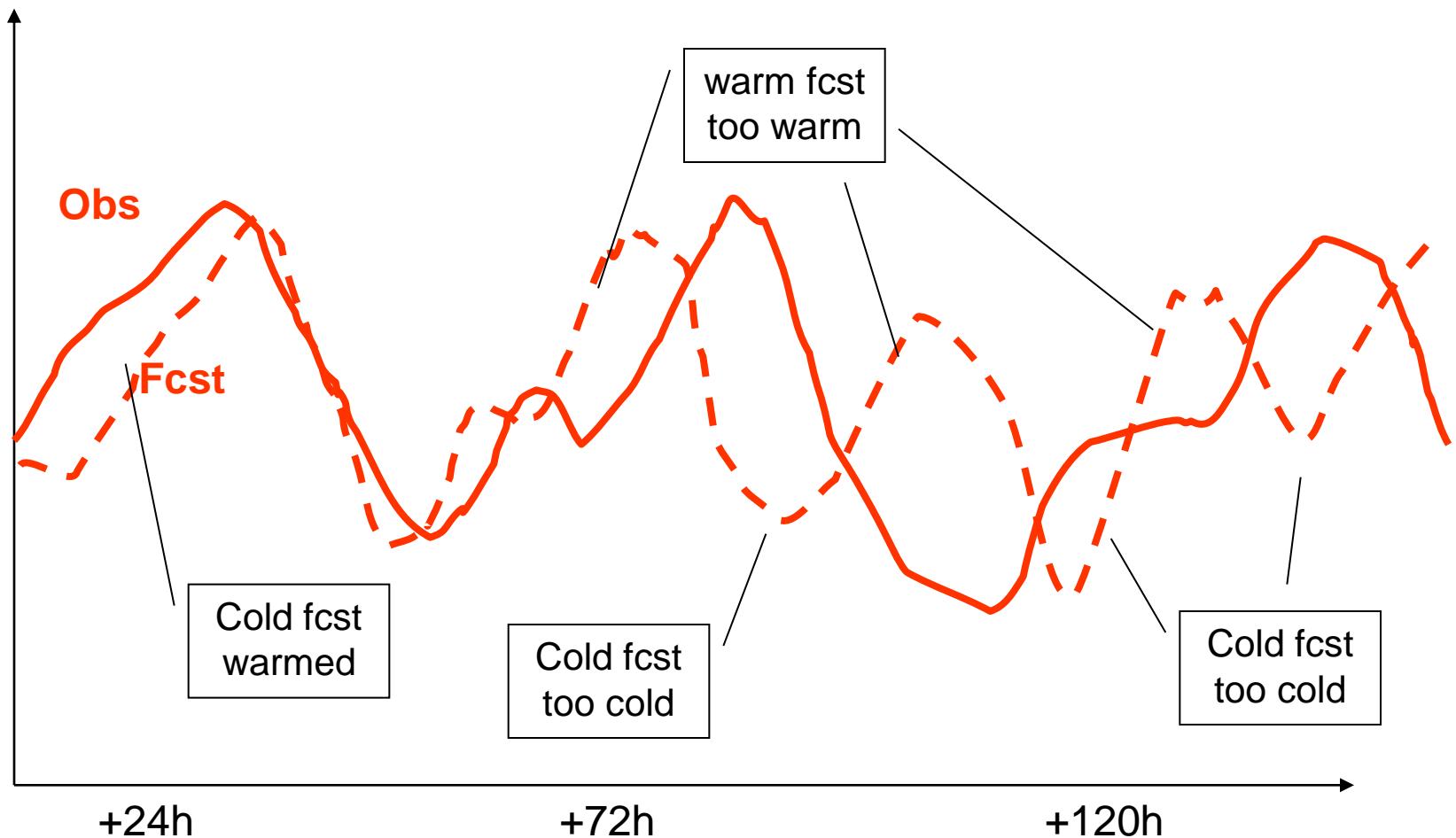
True non-systematic forecast errors (severe phase speed errors)



No **systematic** errors, still warm forecasts are too warm, cold forecasts too cold

Typical medium range forecast (phase and amplitude) errors

Temp (C)



What looks good can be bad – what looks bad can be good

On the interpretation of verification statistics

Decomposition of the RMSE

The complete formula for RMSE

The full mathematical expression for the RMS error (E_j) of a j -day forecast issued on day i verified over N gridpoints over a period of T days

$$E_j = \sqrt{\frac{1}{T} \frac{1}{N} \sum_{n=1}^N \sum_{t=1}^T (f_{i,j} - a_{i+j})^2}$$

But the “Double Penalty Effect” is a manifestation of a more fundamental effect which can be understood through the trivial equation:

$$(a + b)^2 = a^2 + b^2 + 2ab$$

From the RMSE to the MSE

We make things easier for us by considering the *square* of the RMSE

$$E_j^2 = \frac{1}{T} \frac{1}{N} \sum_{n=1}^N \sum_{t=1}^T (f_{i,j} - a_{i+j})^2$$

Simplifying the notations

$$E_j^2 = \frac{1}{T} \frac{1}{N} \sum_{n=1}^N \sum_{t=1}^T (f_{i,j} - a_{i+j})^2$$

The notation is further simplified by replacing the Σ s with an overbar symbolising all temporal and spatial averages. We also skip all the indices.

$$\overline{E^2} = \overline{(f - a)^2}$$

Decomposing the MSE around climate

Decomposing the RMSE

$$E^2 = \overline{(f - a)^2}$$

Introduce c as the climate value of the verifying day

$$E^2 = \overline{(f - c + c - a)^2}$$

Reposition c to form $f-c$ and $a-c$

$$E^2 = \overline{((f - c) - (a - c))^2}$$

Apply $(a+b)^2 = a^2 + b^2 + 2ab$

$$E^2 = \overline{(f - c)^2} + \overline{(a - c)^2} - 2\overline{(f - c)(a - c)}$$

Each of these three terms has its own story to tell

What looks bad might be good...

The *improvement* of the model, as a simulation of the atmospheric system, may therefore appear as *deterioration* of the quality of the model!

An increase in forecast variability increases A_f^2

...compensates the decrease of the RMSE due to improved forecasts

$$E^2 = \overline{(f - c)^2} + \overline{(a - c)^2} - 2\overline{(f - c)(a - c)}$$

...to the level of the observed variability A_a^2

The Error Saturation Level

The atmosphere's variability

When $f=c$ the first and last terms disappear

$$E^2 = \overline{(f - c)^2} + \overline{(a - c)^2} - 2\overline{(f - c)(a - c)}$$

$$E^2 \rightarrow \overline{(a - c)^2}$$

$$E^2 \rightarrow A_a^2$$

We take the square root....

$$E \rightarrow A_a$$

Which is the error level for a purely climatological statement

The RMS error saturation level

When the forecasts start to lose skill and the RMSE start to approach high error levels the last term disappears

$$E^2 = \overline{(f - c)^2} + \overline{(a - c)^2} - 2\overline{(f - c)(a - c)}$$

$$E^2 \rightarrow A_f^2 + A_a^2$$

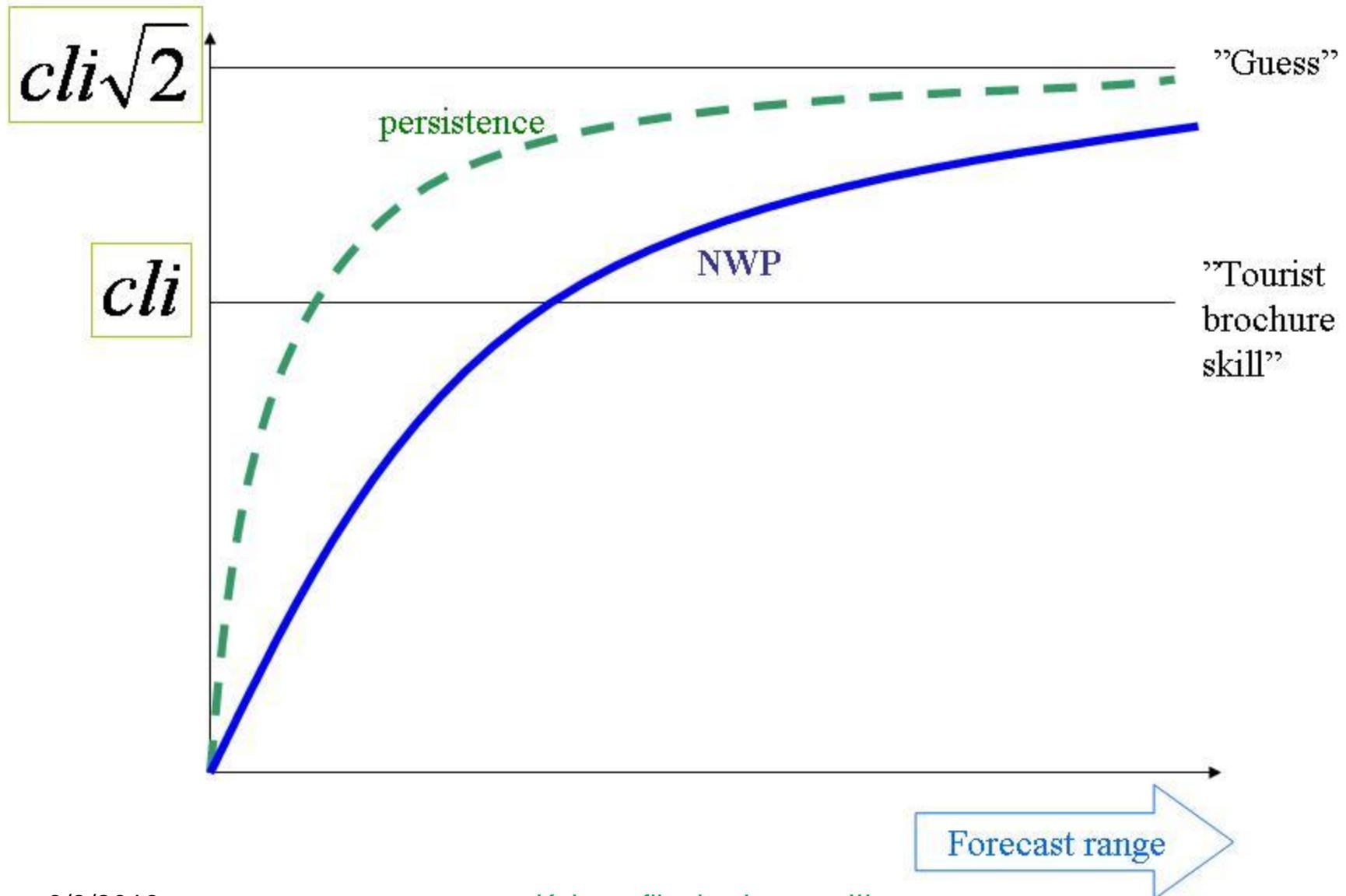
$$E^2 \rightarrow 2A_a^2$$

$$E \rightarrow A_a \sqrt{2}$$

Which is the uppermost error level for a realistic NWP model, 41% above the error level of a purely climatological statement

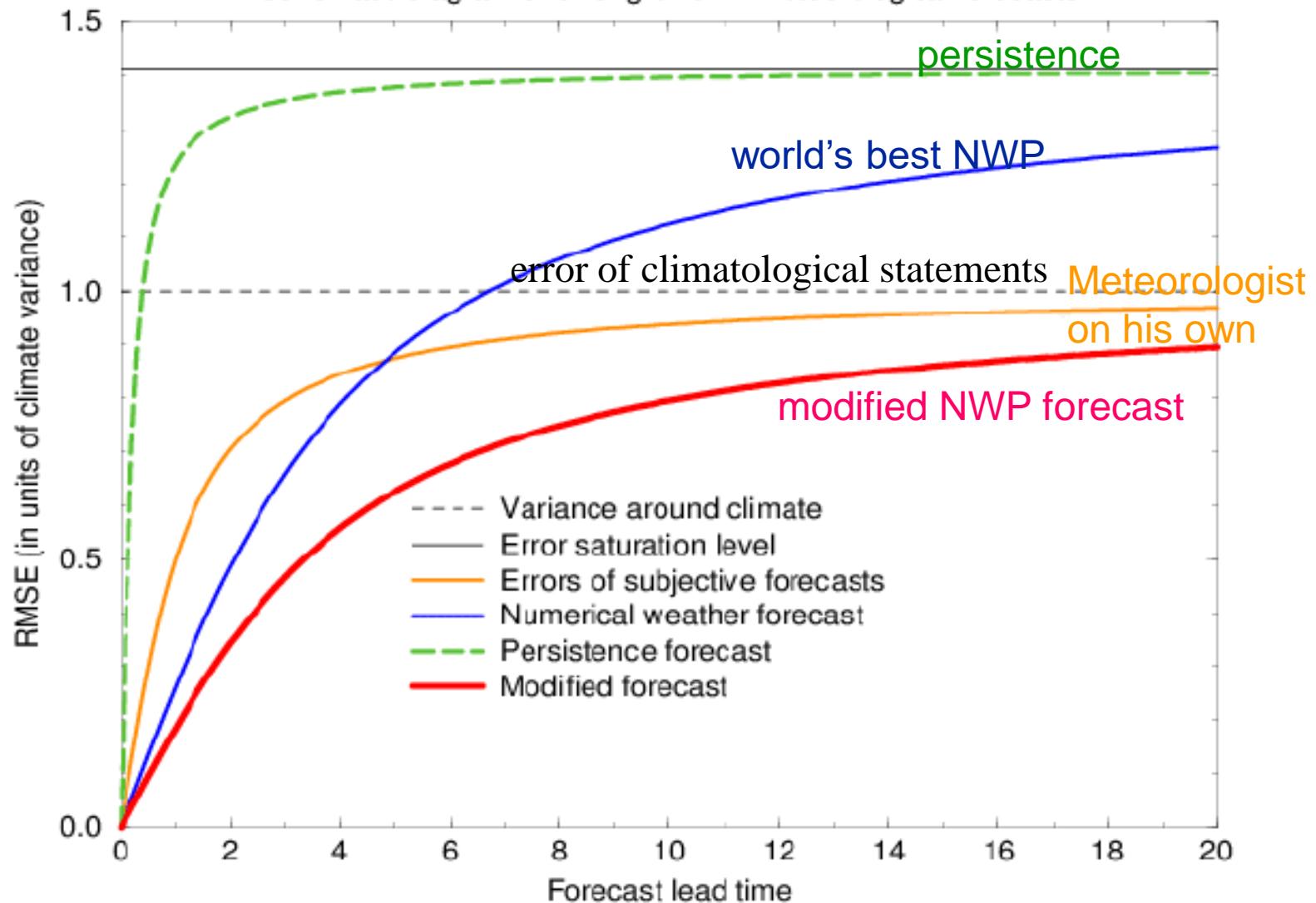
It is also called *The Error Saturation Level (ESL)*

Forecast Error Growth



Forecast error growth and saturation levels

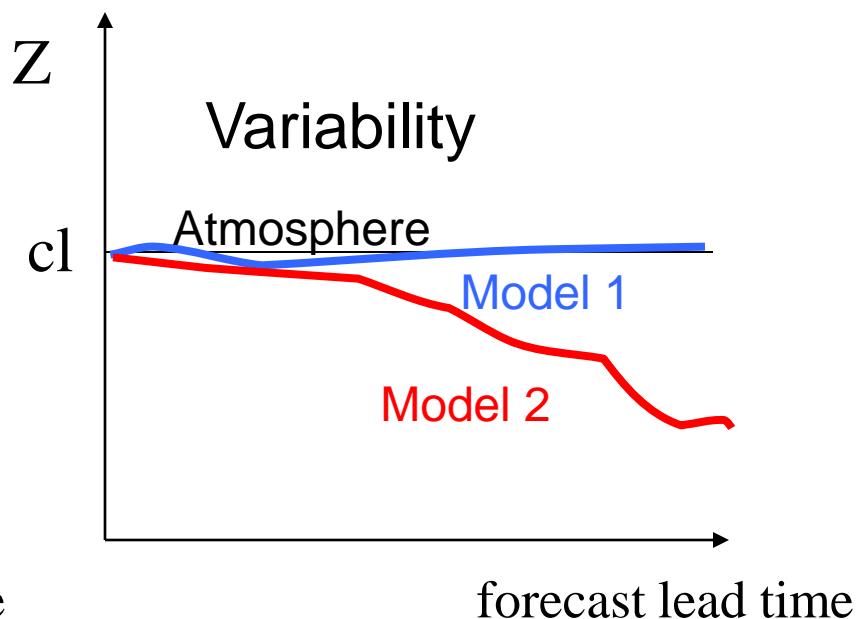
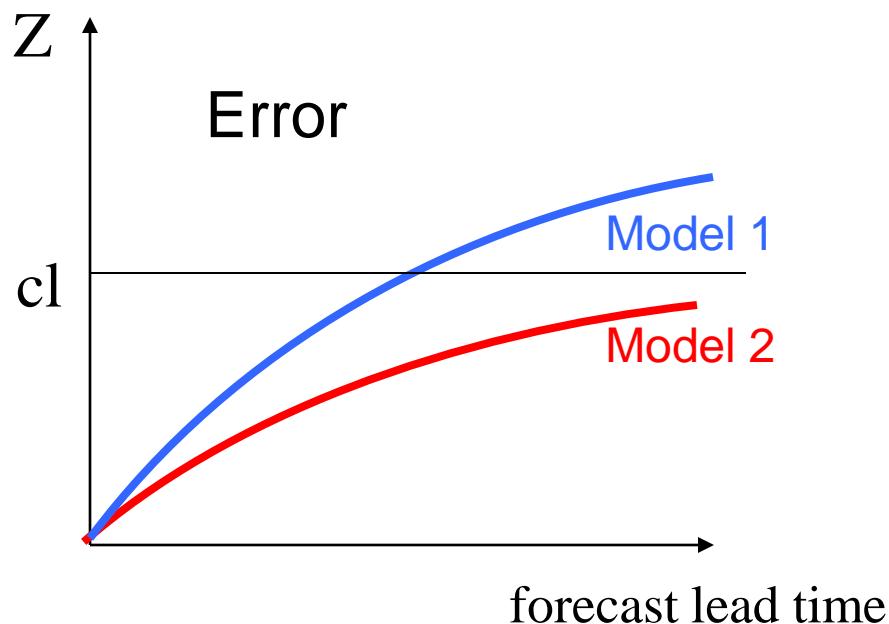
Schematic diagram of error growth in meteorological forecasts



The use of the decomposition of the RMSE

$$E^2 = \overline{(f-a)^2} = \overline{(f-c)^2} + \overline{(a-c)^2} - 2\overline{(f-c)(a-c)}$$

model variab. atm. variab. “skill”

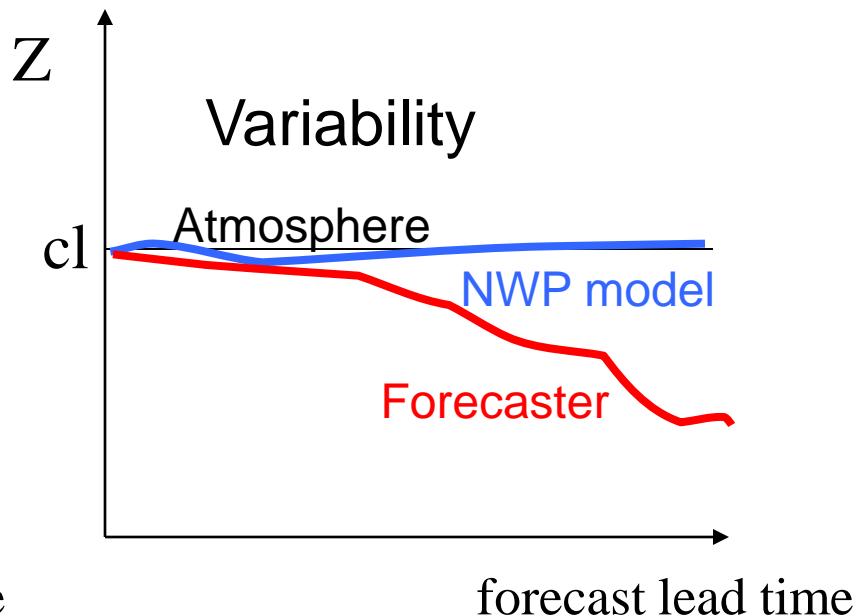
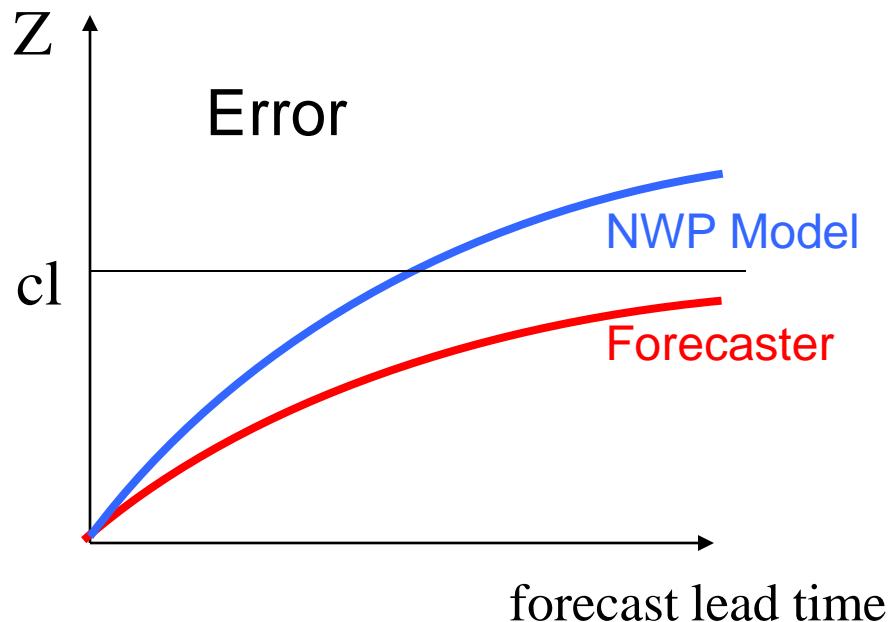


The decrease in variability of Model 2 (bad) gave low (good) RMSE verifications

But the same does not apply to a forecaster

$$E^2 = \overline{(f - a)^2} = \overline{(f - c)^2 + (a - c)^2} - 2\overline{(f - c)(a - c)}$$

Forecast variab. Atm. variab. "skill"

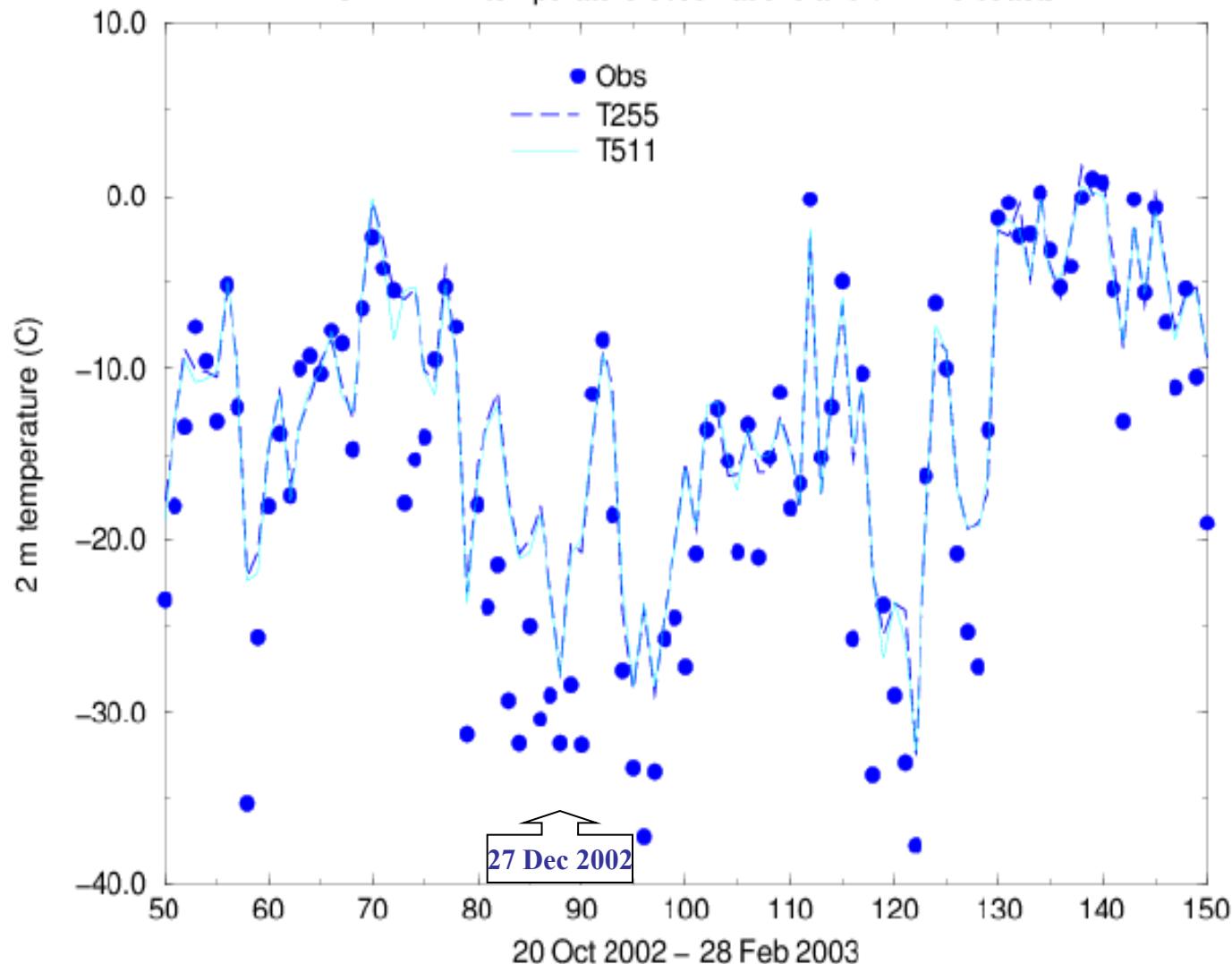


The decrease in variability of the good forecaster gave low (good) RMSE verifications

What's the point of compensating
for under variability if it does not
improve the forecasts sufficiently?

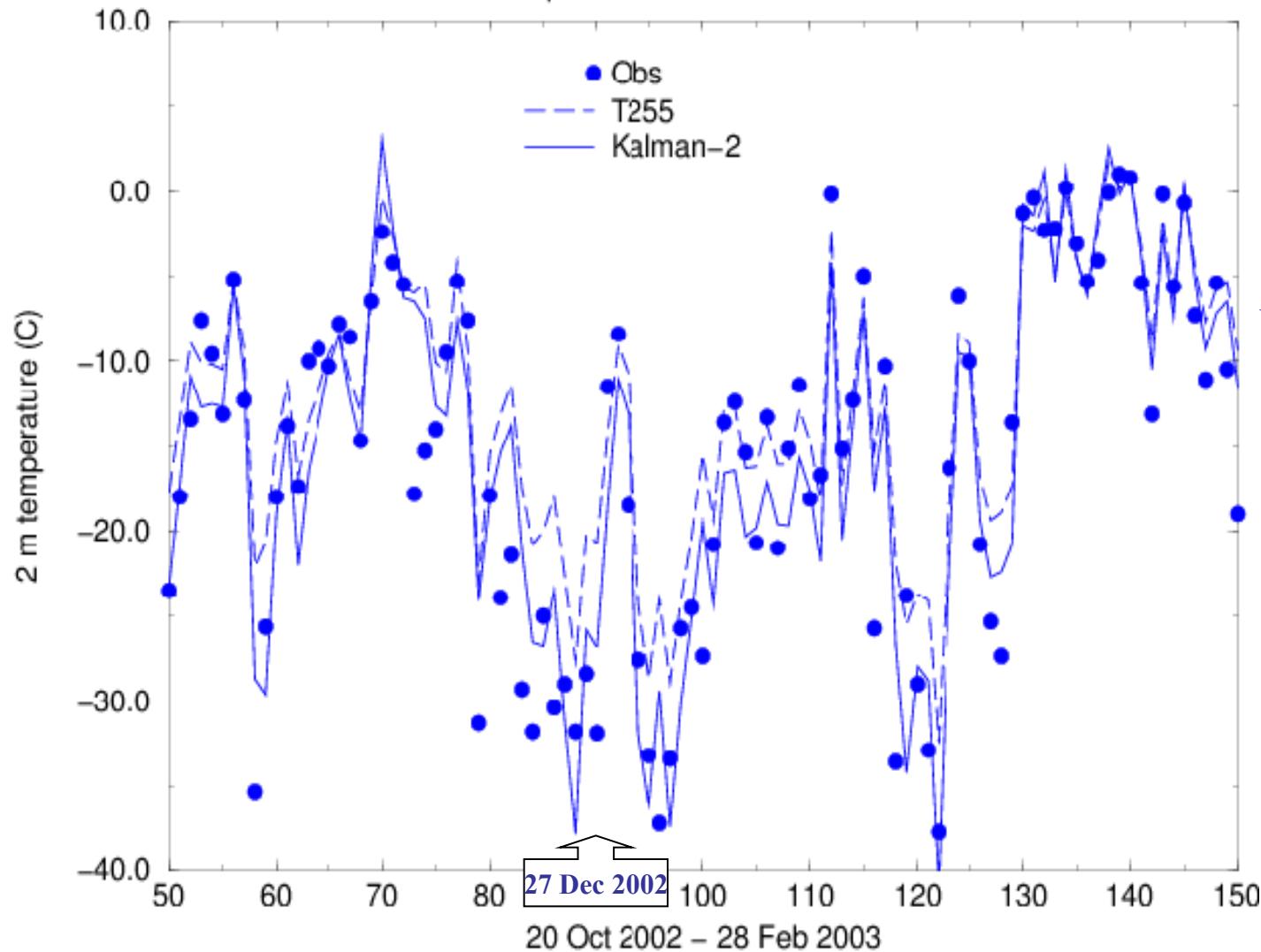
Winter temperatures in Sodankyla, N Finland

ECMWF 2 m temperature observations and +12h forecasts



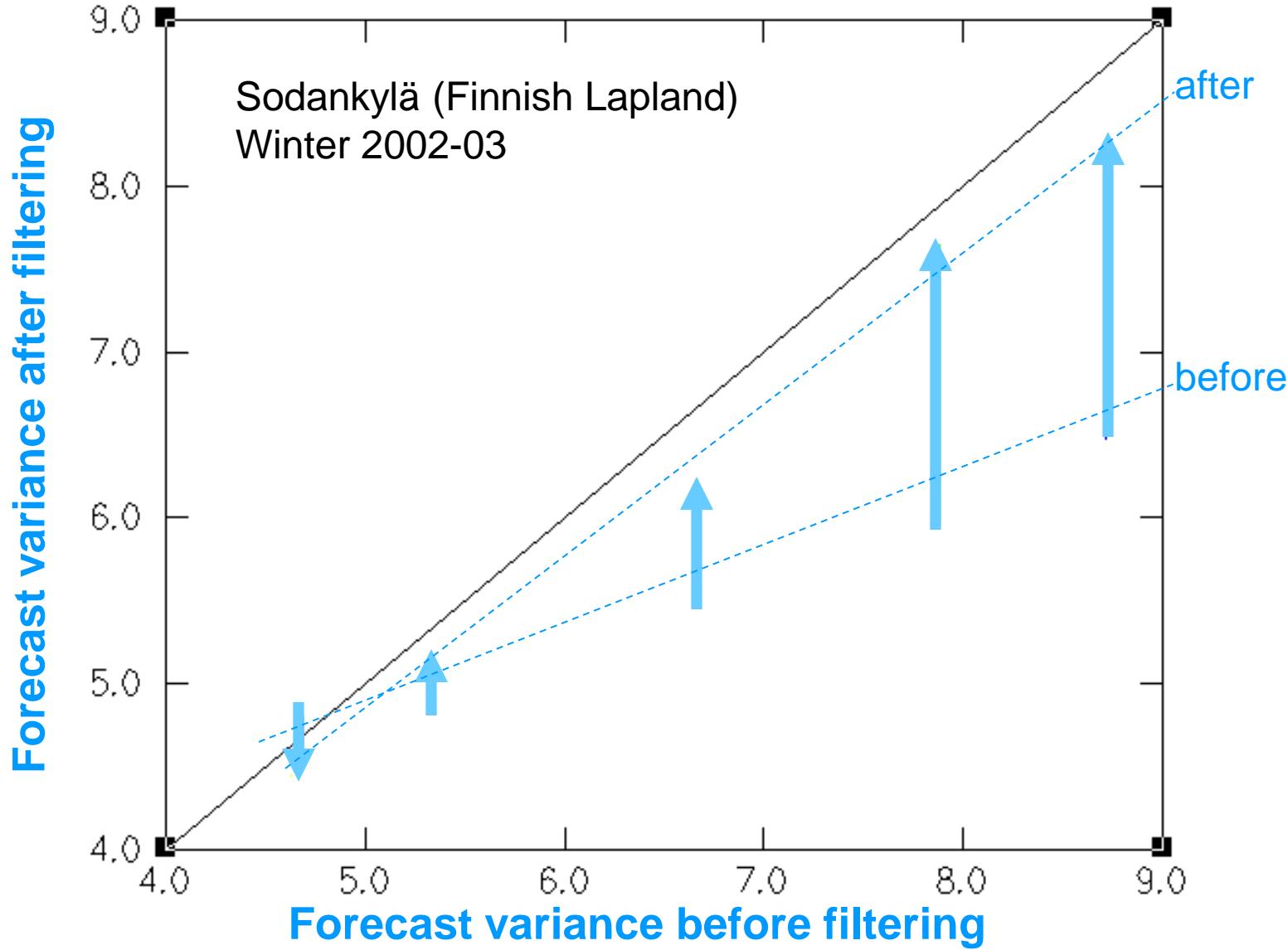
Winter temperatures in Sodankyla, N Finland

ECMWF 2 m temperature observations and +12h forecasts

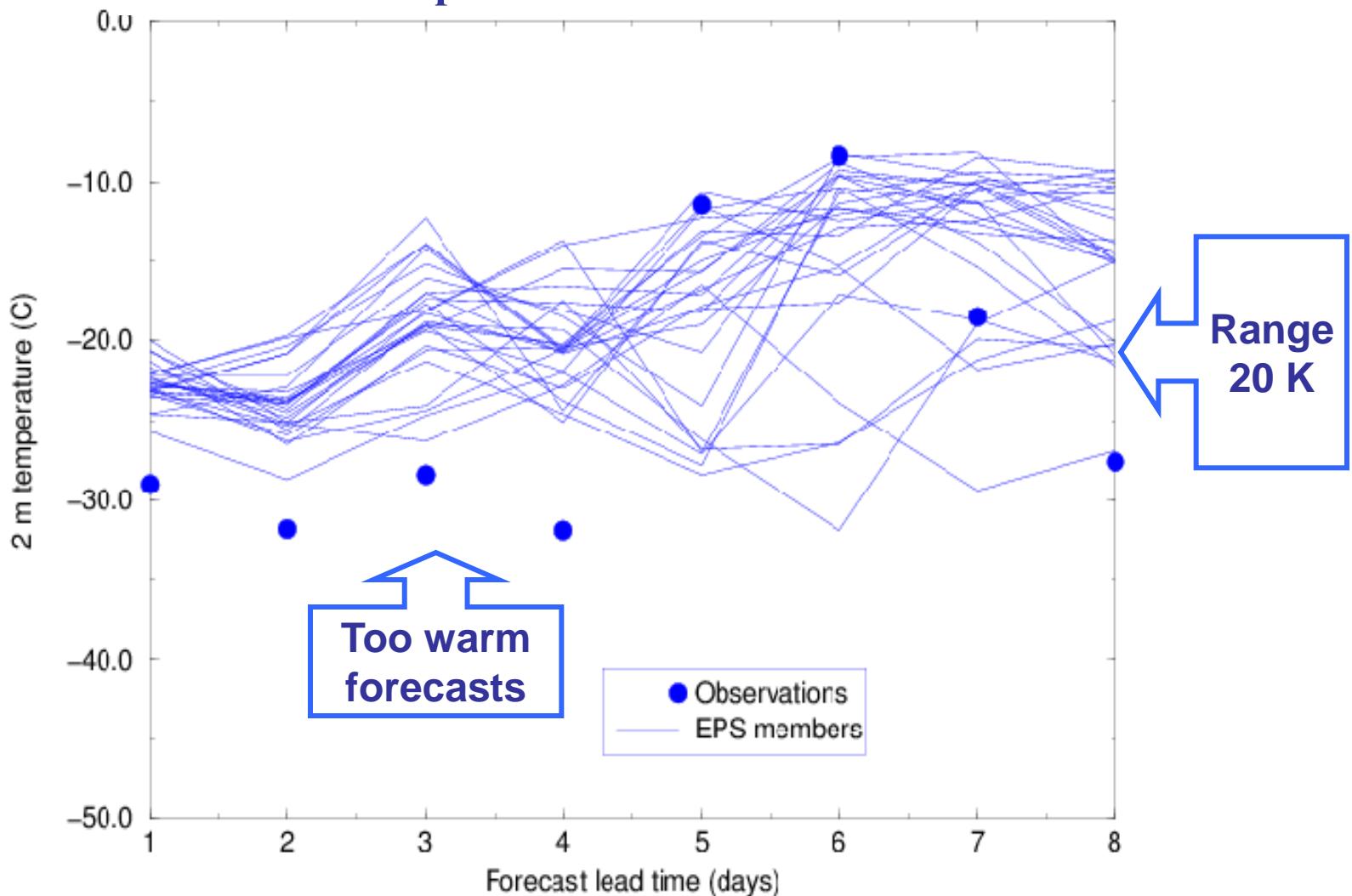


The mean error (ME) decreased by 1-2 K but the RMSE only by 0.5 K
The solution lies in the ensemble approach...

The Kalman 2 filter improvement of forecast variance

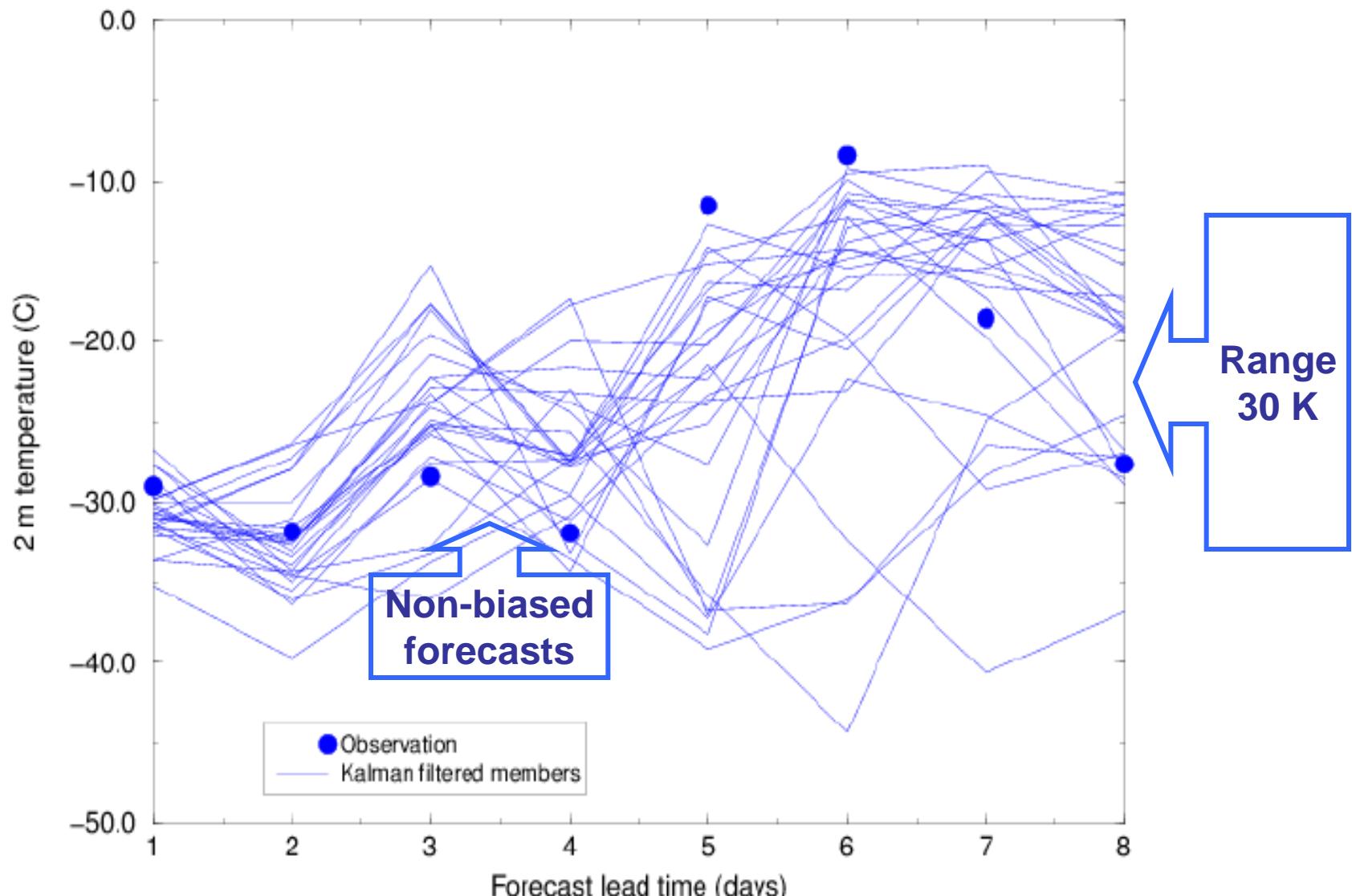


The T255 (and T511) have problems with temperatures below -25 °C

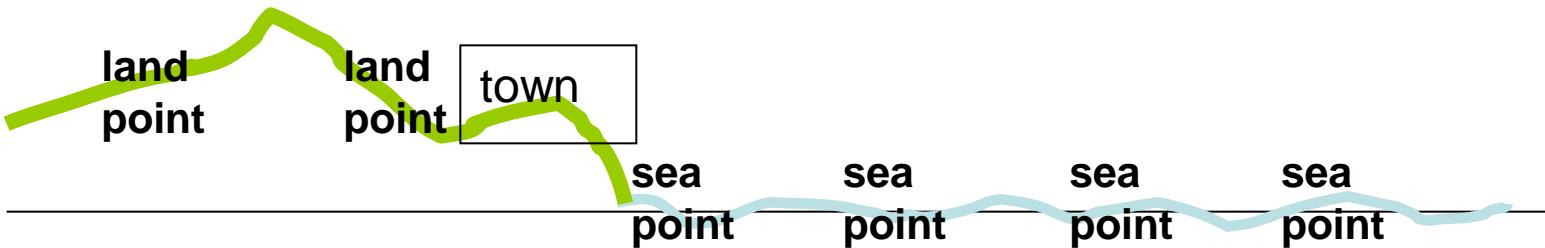


2 m temperature ensemble forecast for Sodankyla

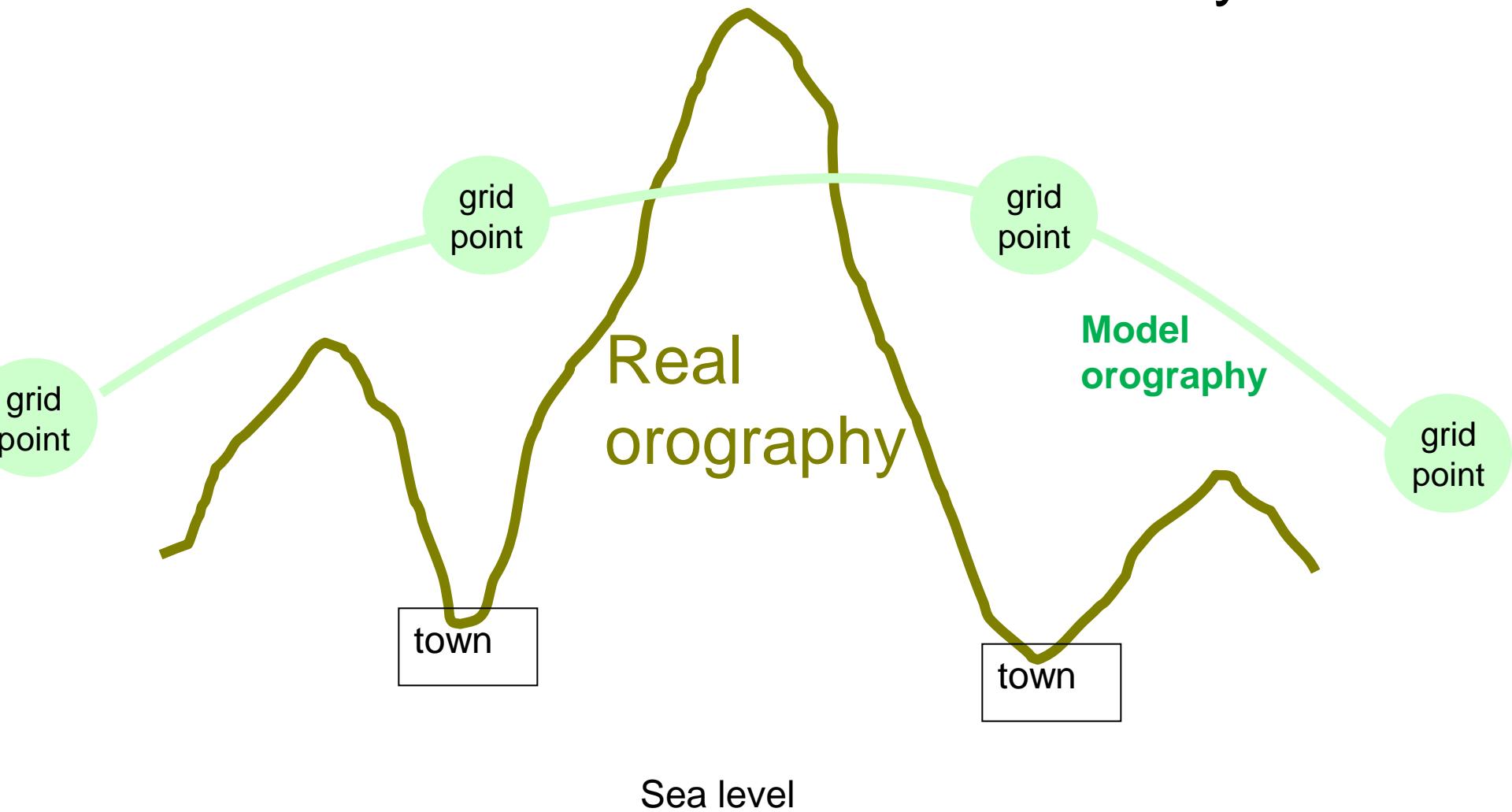
Kalman filtered EPS forecast 27 Dec 2002



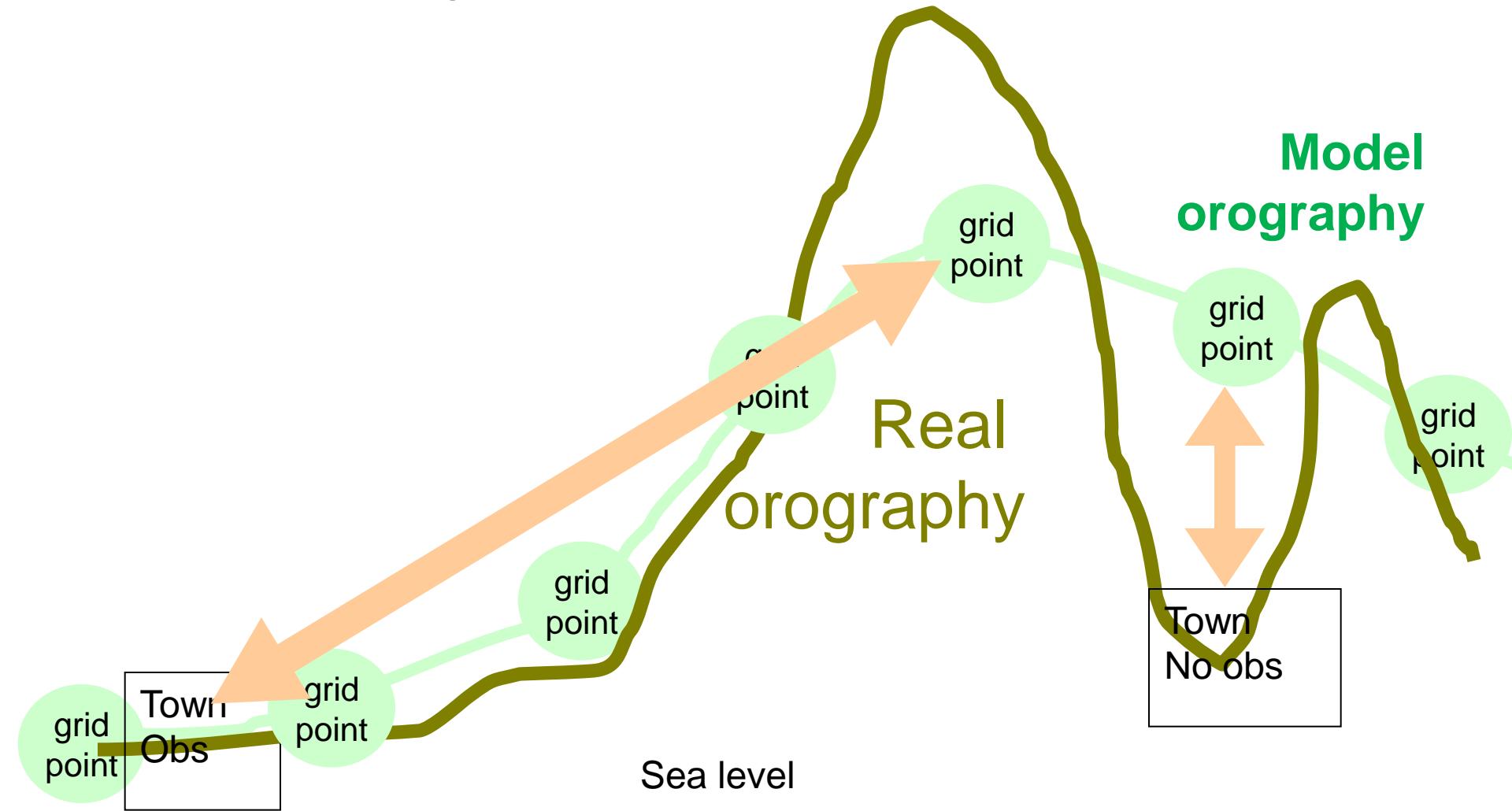
The issue so far has been the relation between observations and grid points at about the same height



A more difficult problem is to adjust model level data to stations in valleys

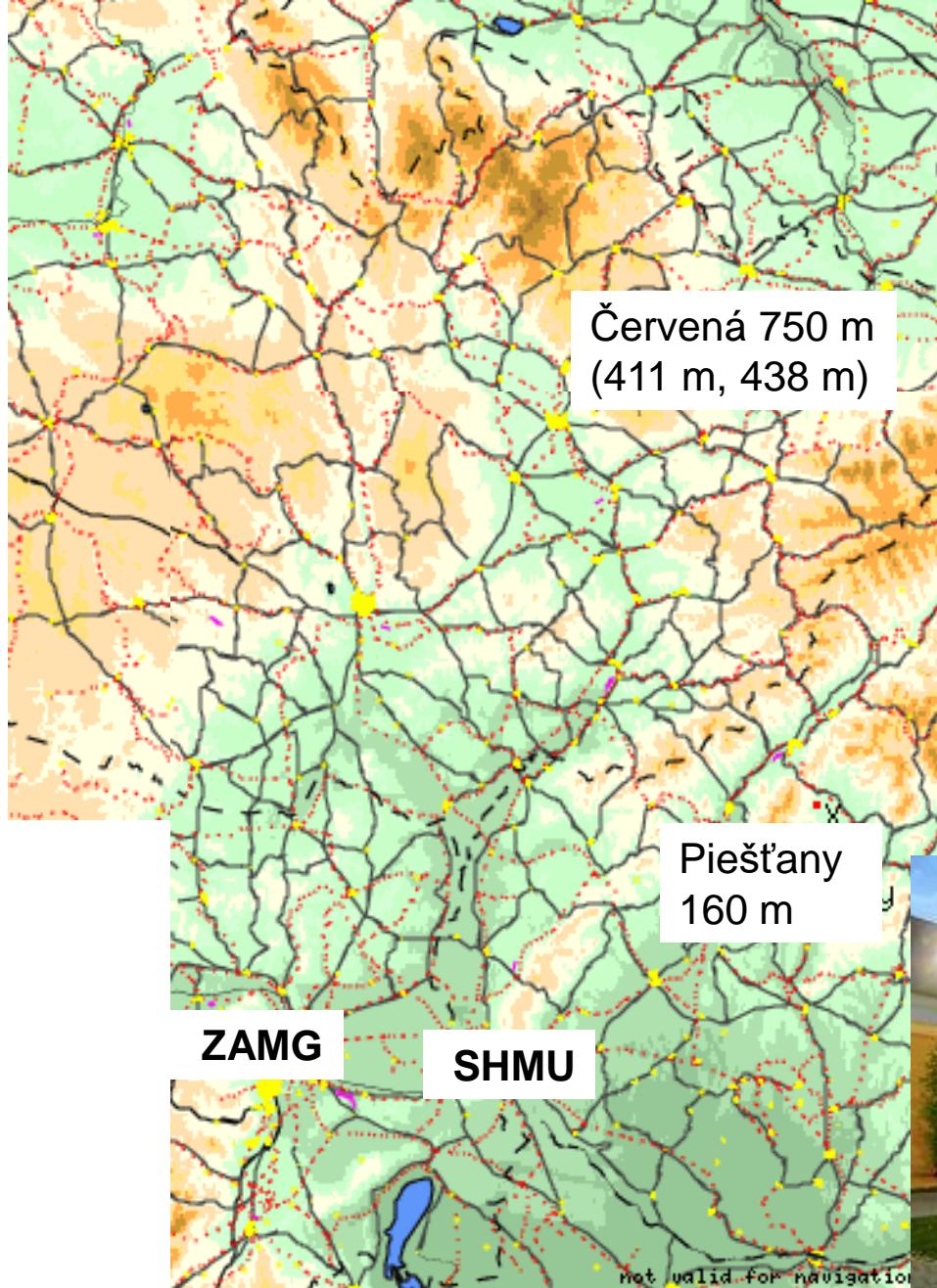


For practical reason my comparison between high altitude grid point data low altitude observations also involved great horizontal differences



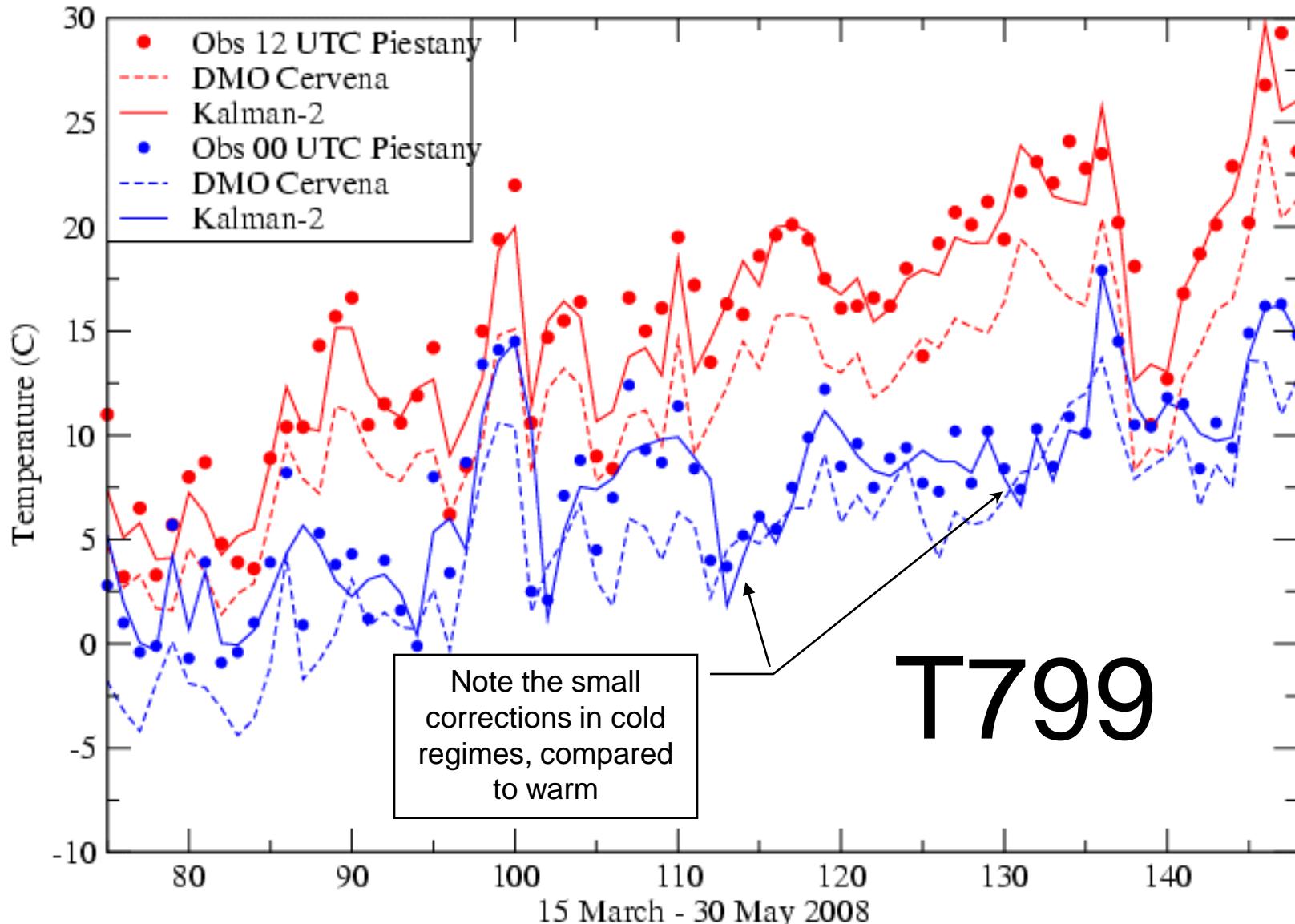
Comparison between high altitude grid points and low altitude observations, which for practical reasons also happened to be situated at some considerable distance





Kalmanfiltering ECMWF D+1 forecasts for Piestany (160 m)

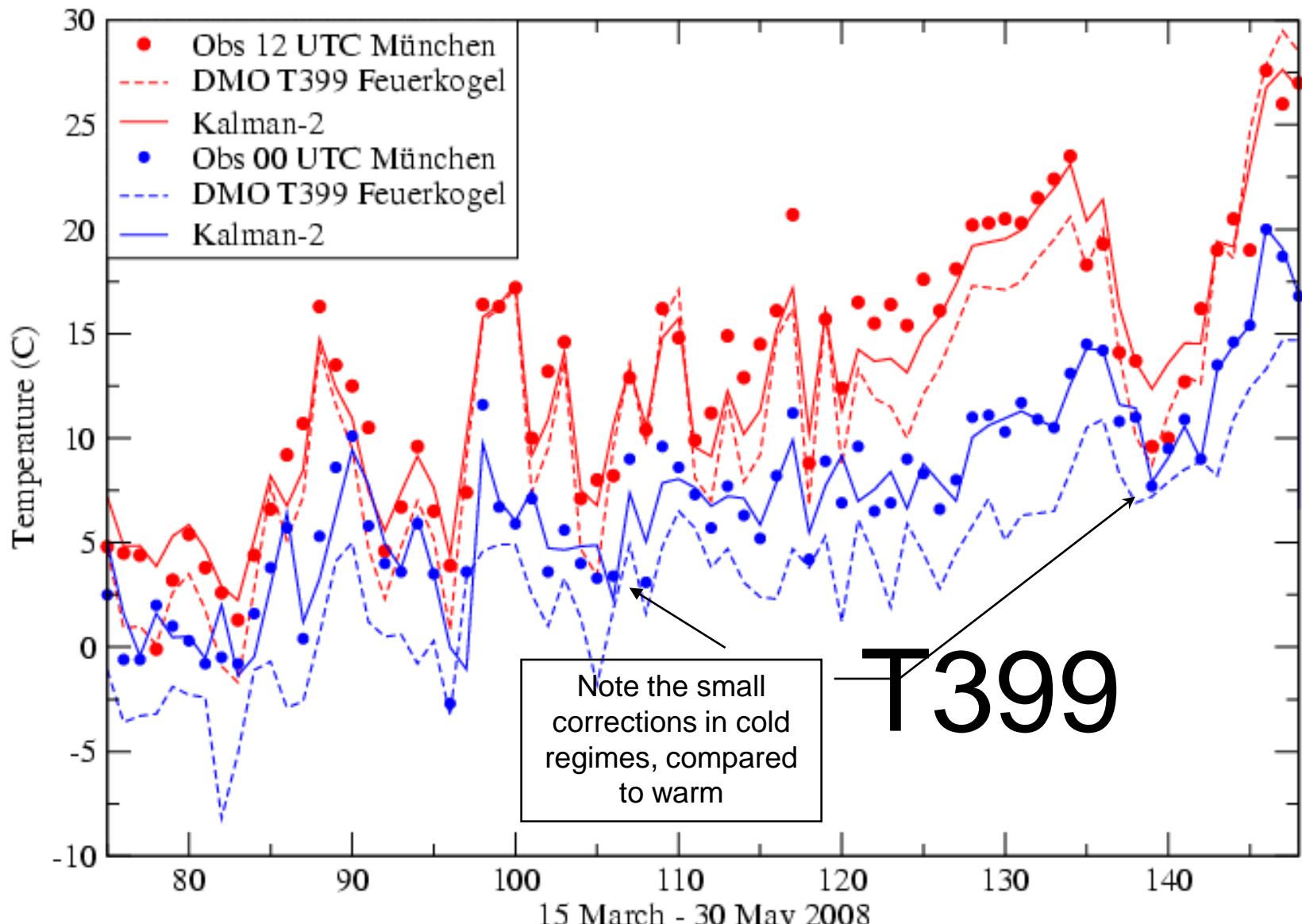
...against ECMWF D+1 forecasts for Cervena 354 m (model height)





Kalmanfiltering ECMWF D+1 forecasts for München (447 m)

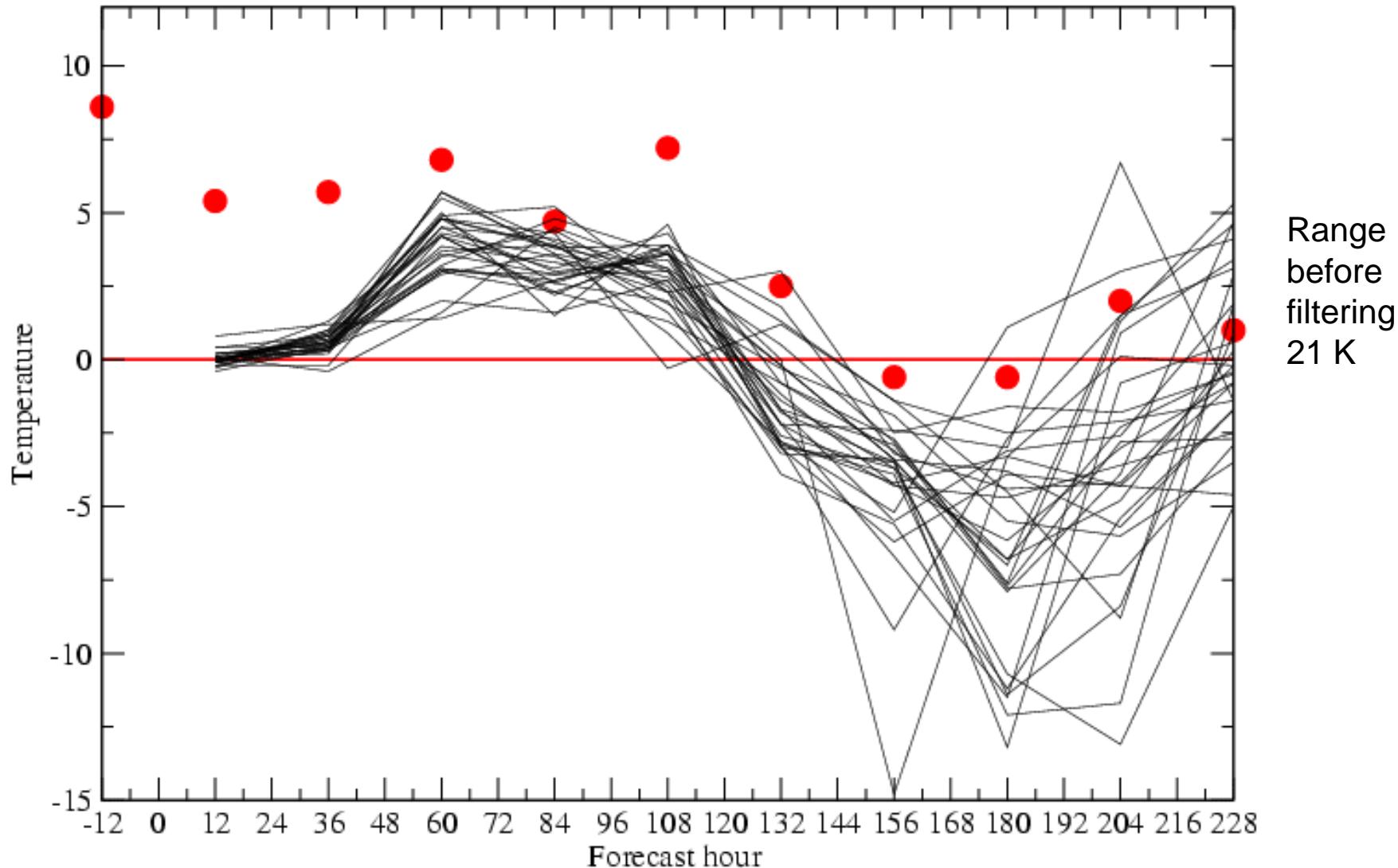
...against ECMWF D+1 EPS Control forecasts for Feuerkogeln 1362 m (model height)



... or the Central European mountain sites:

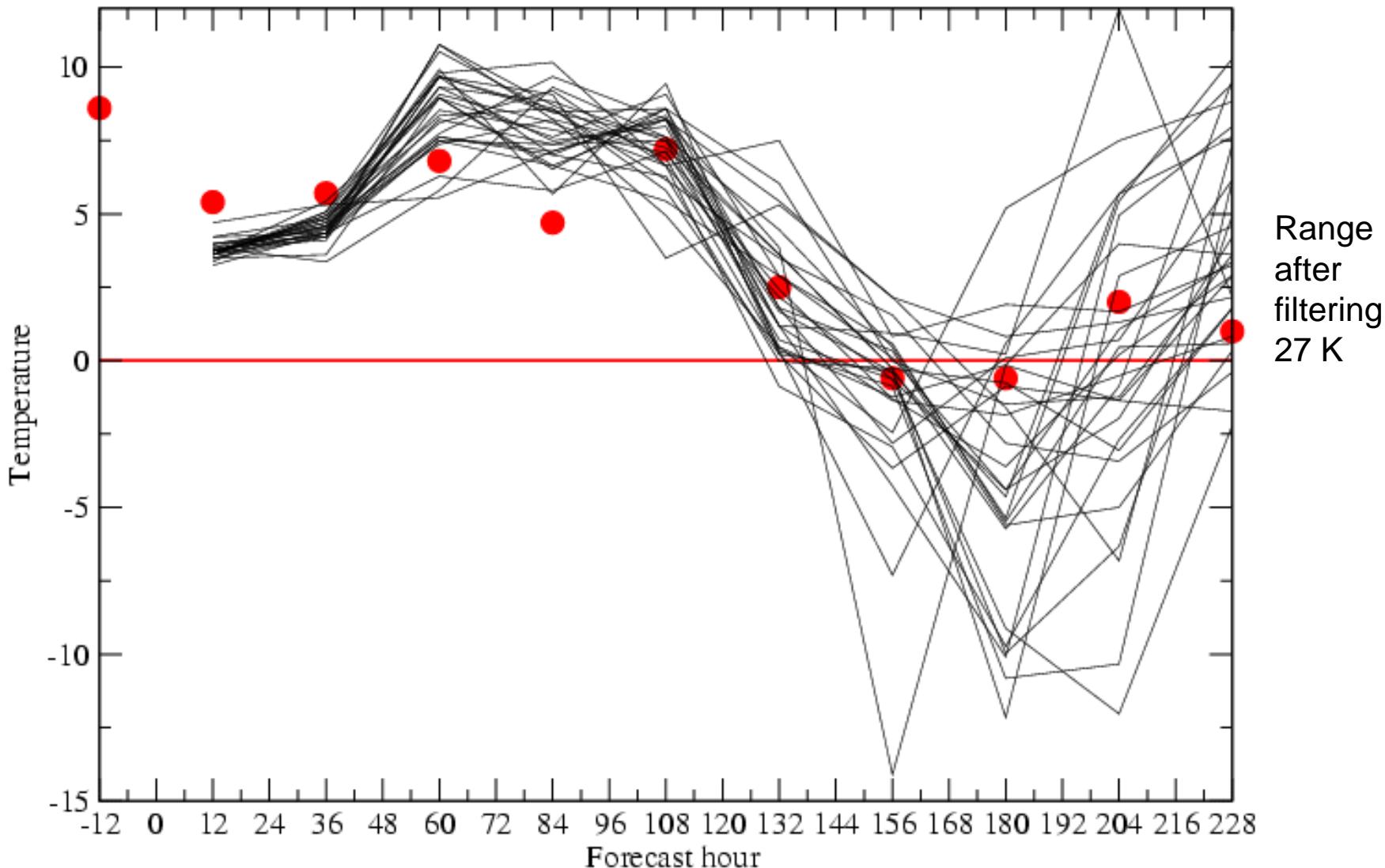
Forecast for München, based on forecast for Feuerkogel before Kalman filtering

The 12 March 2008 12 UTC Ensemble forecast for 2 m temperature and verifying observations



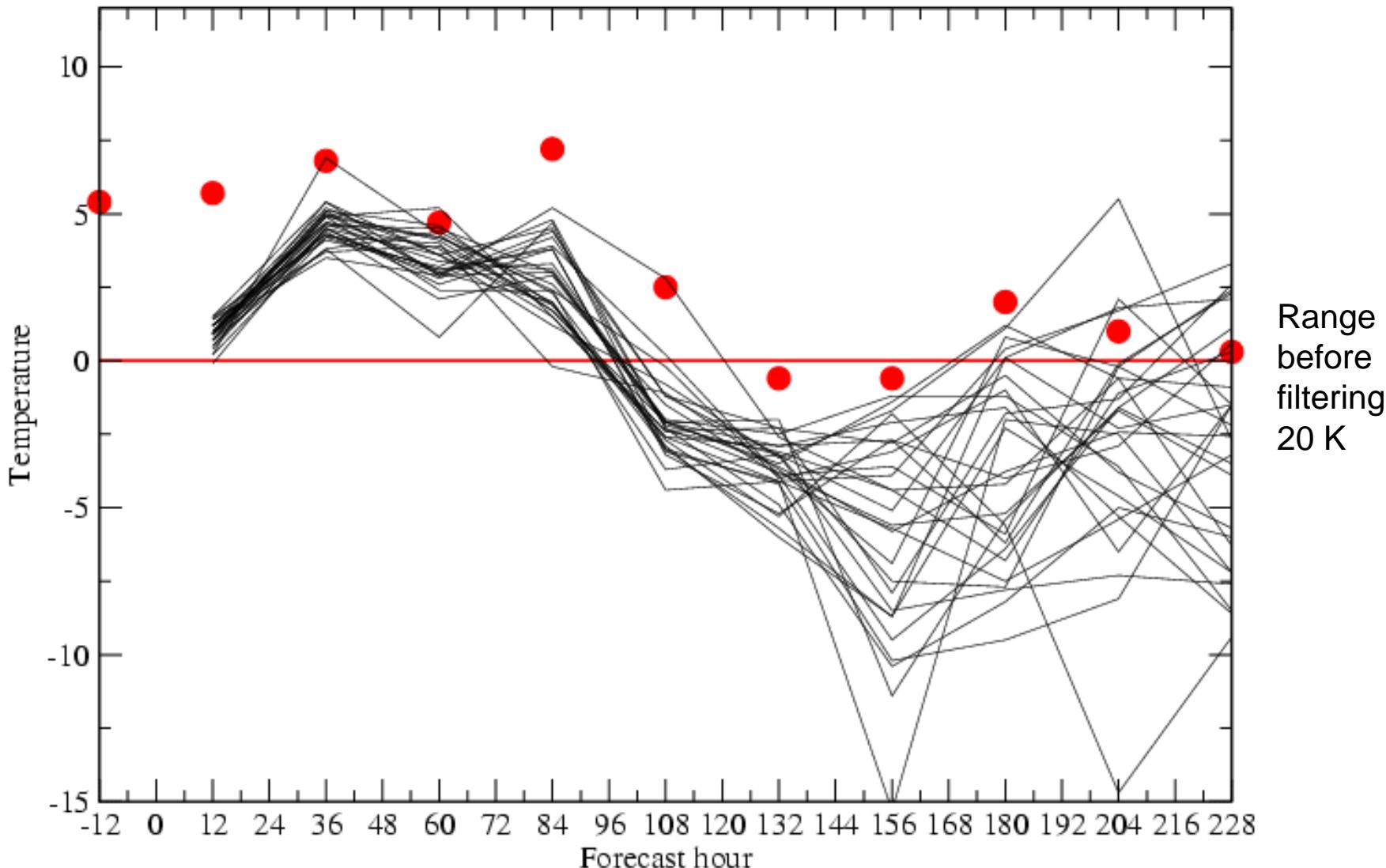
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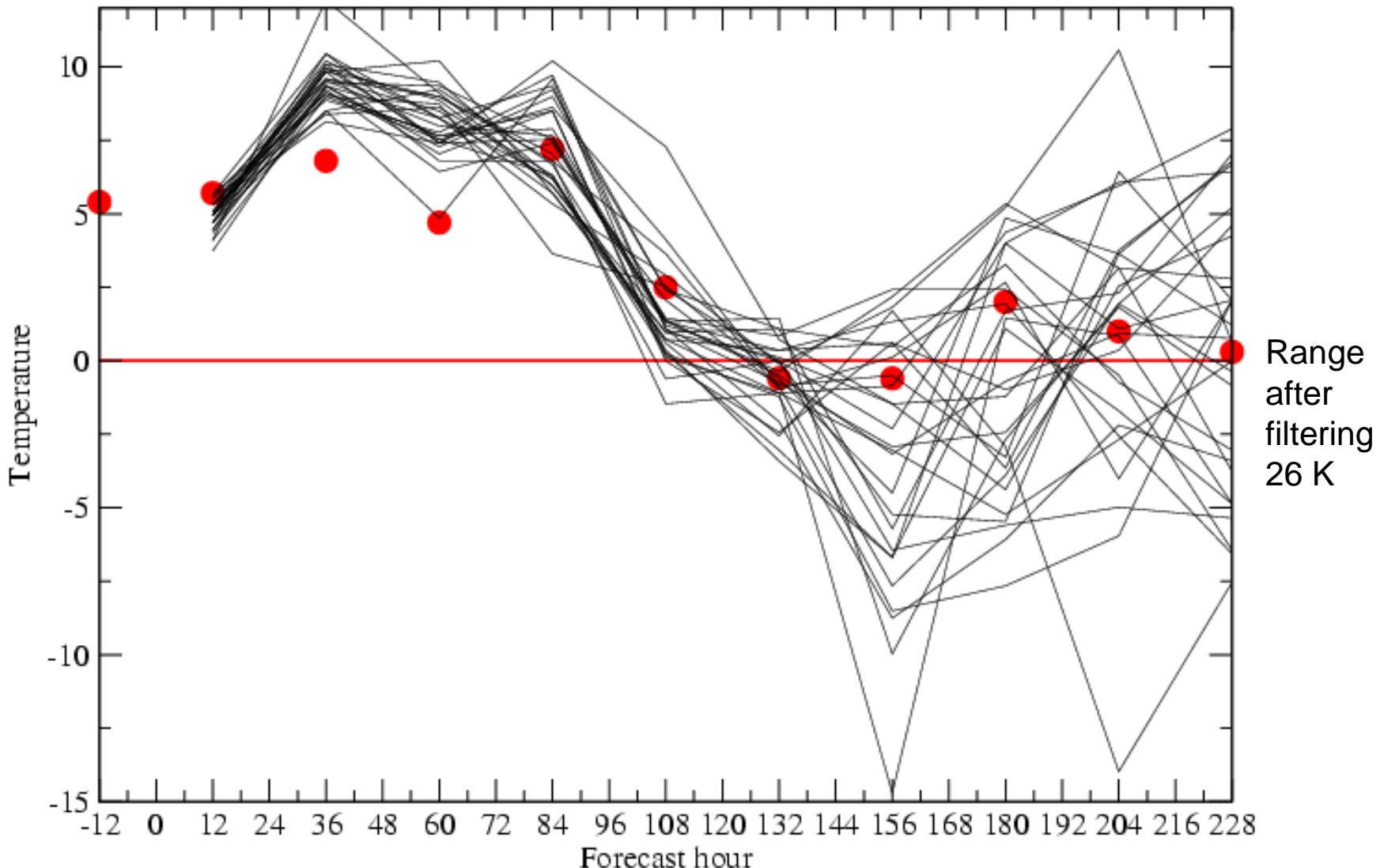
Forecast for München, based on forecast for Feuerkogel before Kalman filtering

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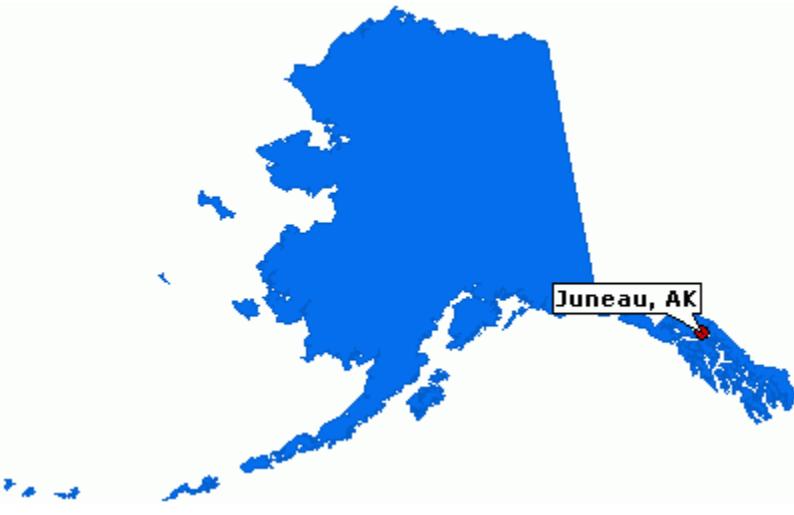
Forecast for München, based on forecast for Feuerkogel after Kalman filtering

The 13 March 2008 12 UTC Ensemble forecast for 2 m temperature and verifying observations



Two Kalman filters operating in parallel

The case of Juneau in Alaska



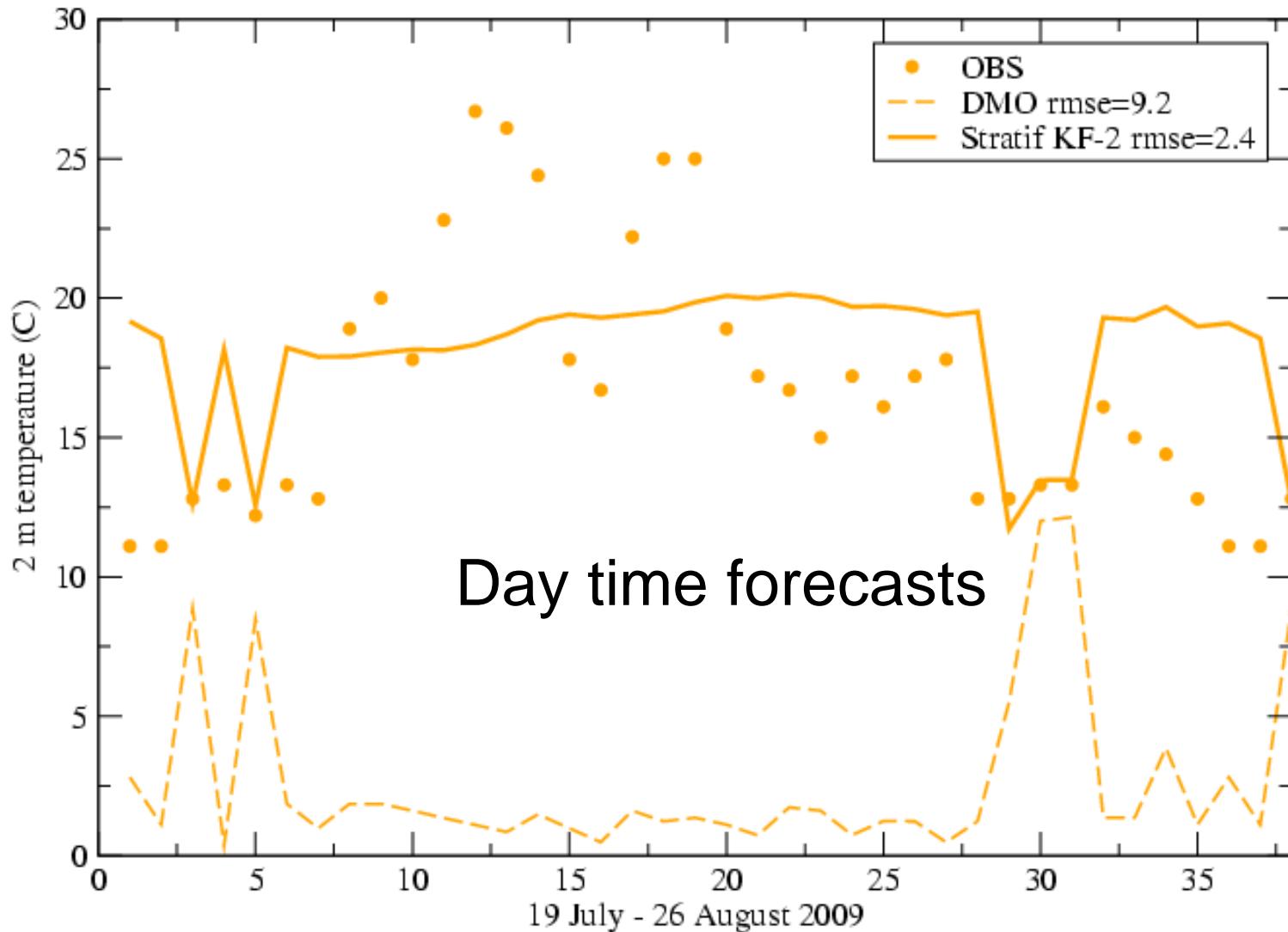
Wind from NE gave NWP $\approx 2^{\circ}\text{C}$
Wind from SE gave NWP $\approx 10^{\circ}\text{C}$

One filter only worked for
NWP $0\text{-}5^{\circ}\text{C}$, the other
one only for $5\text{-}15^{\circ}\text{C}$



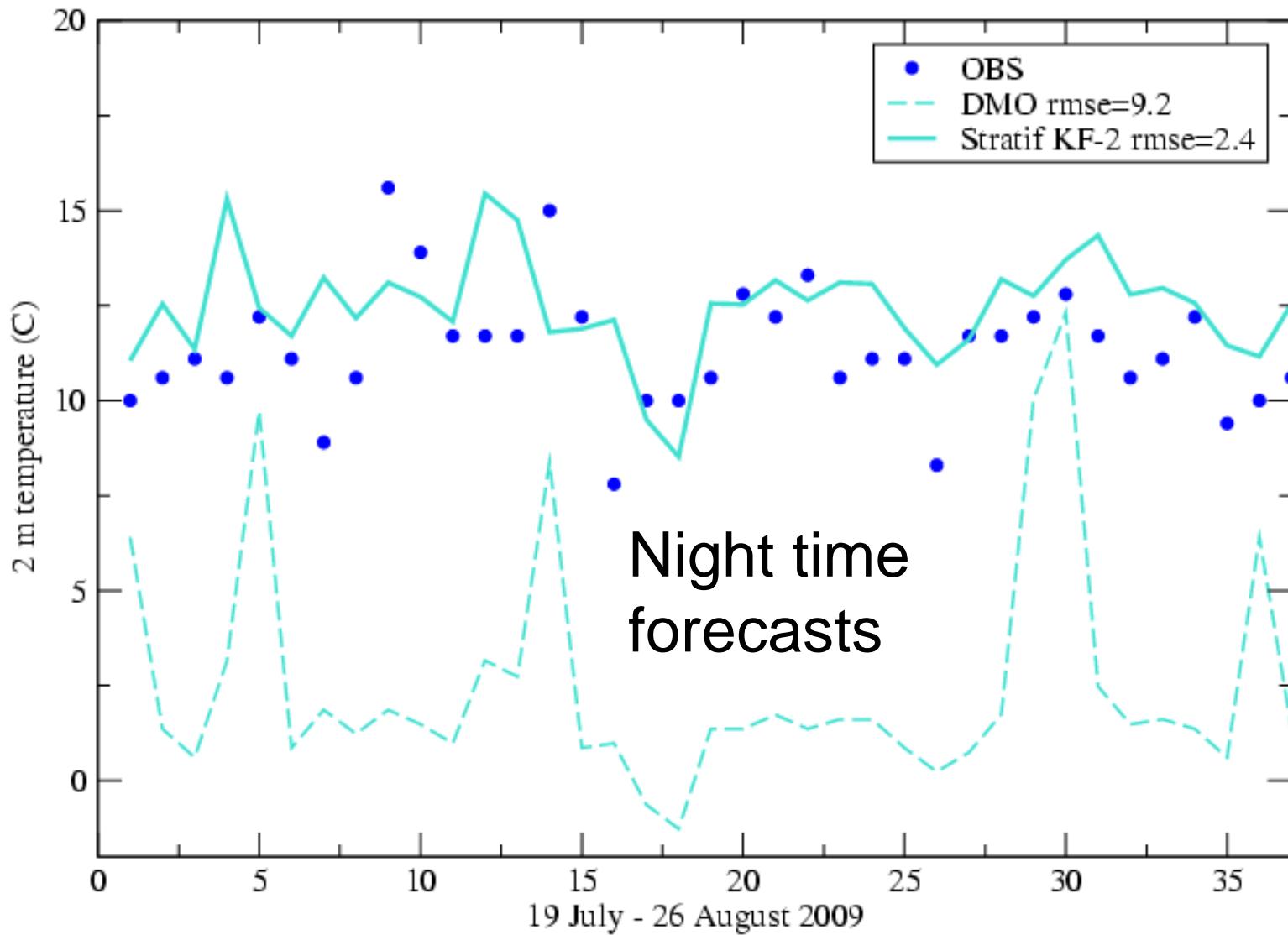
Kalman-2 filtering the UKMO global model

00 UTC + 24h 2 m temperature forecast for Juneau (SE Alaska)



Kalman-2 filtering the UKMO global model

12 UTC + 24h 2 m temperature forecast for Juneau (SE Alaska)



The reason for the bad forecasts for Juneau in Alaska?

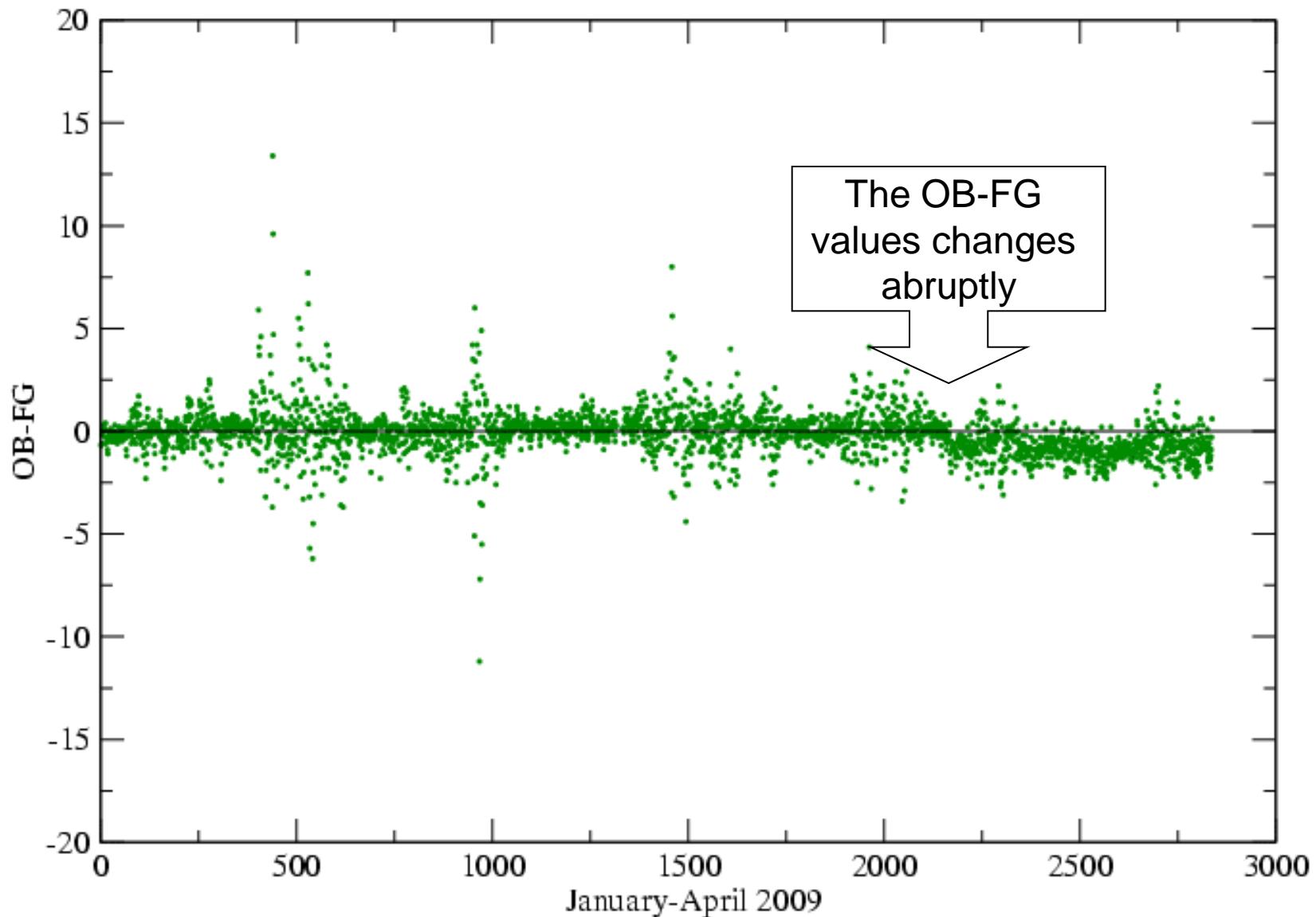


The model had kept ice in the region
From the last Ice Age?

Other simple applications of Kalman filtering

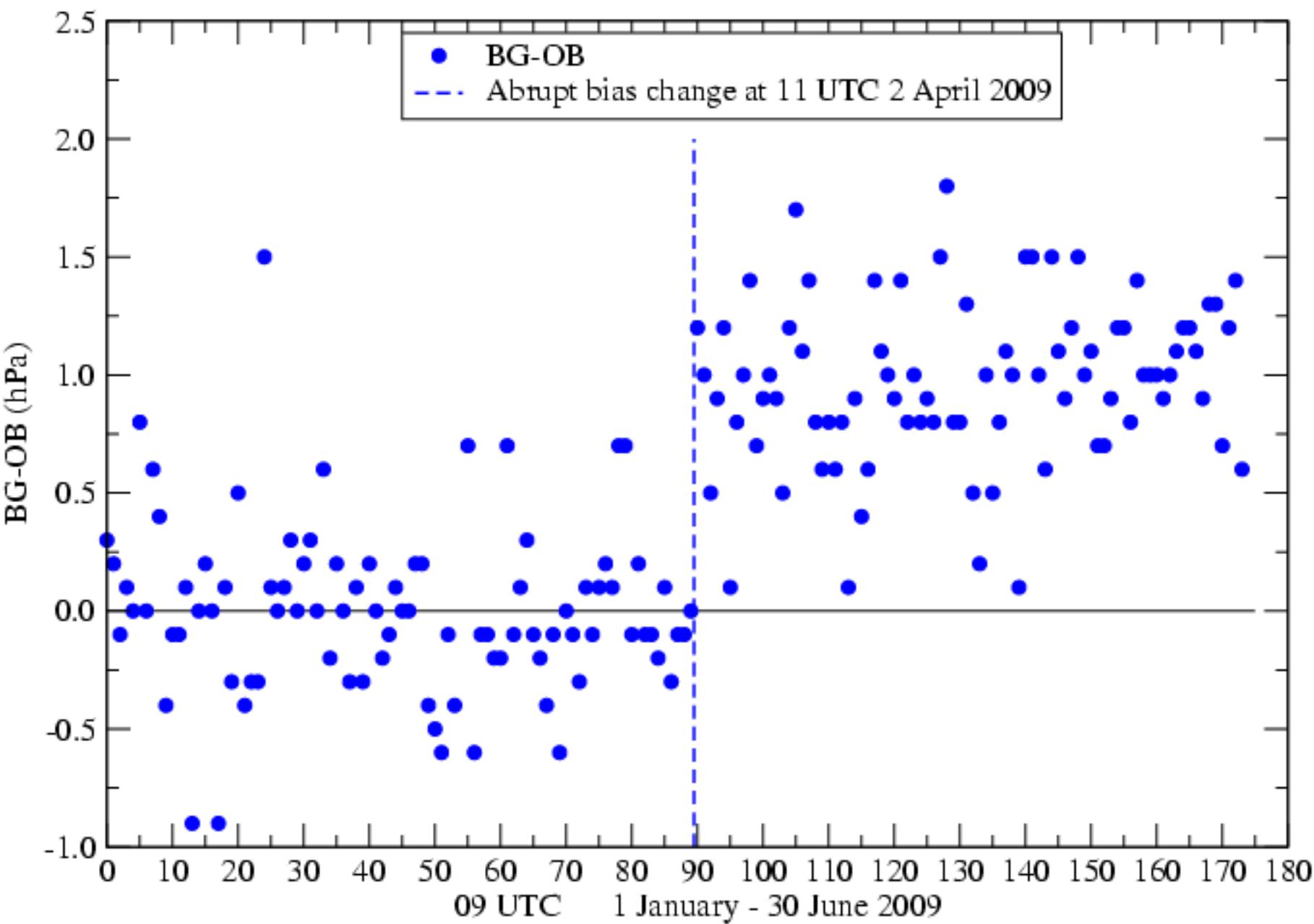
Kalman-1 filtering of biased platforms

OB-FG observations from 03866



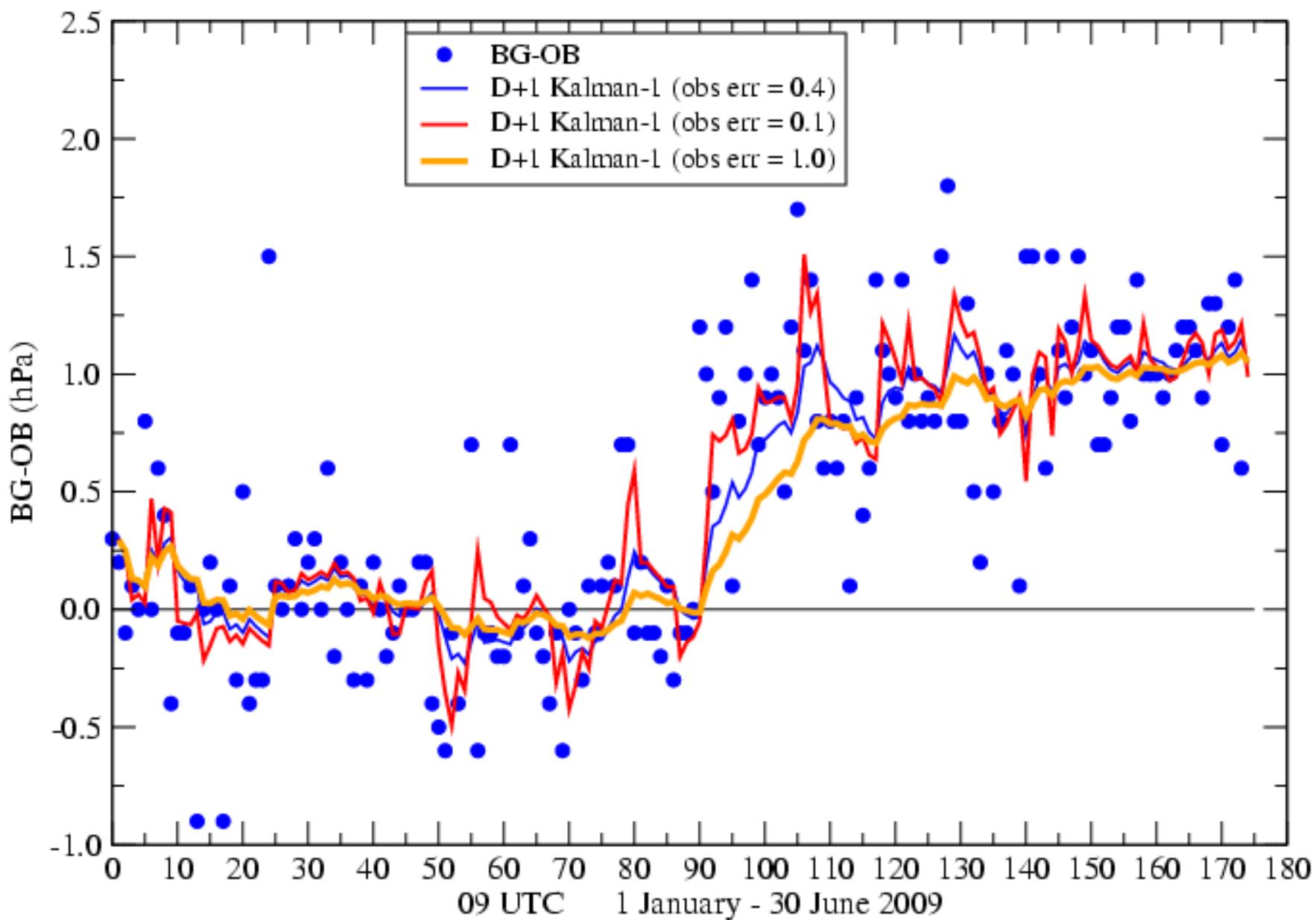
Kalman filtering of MSLP biases

SYNOP 03866, St Cathrine's Point, Isle of Wight



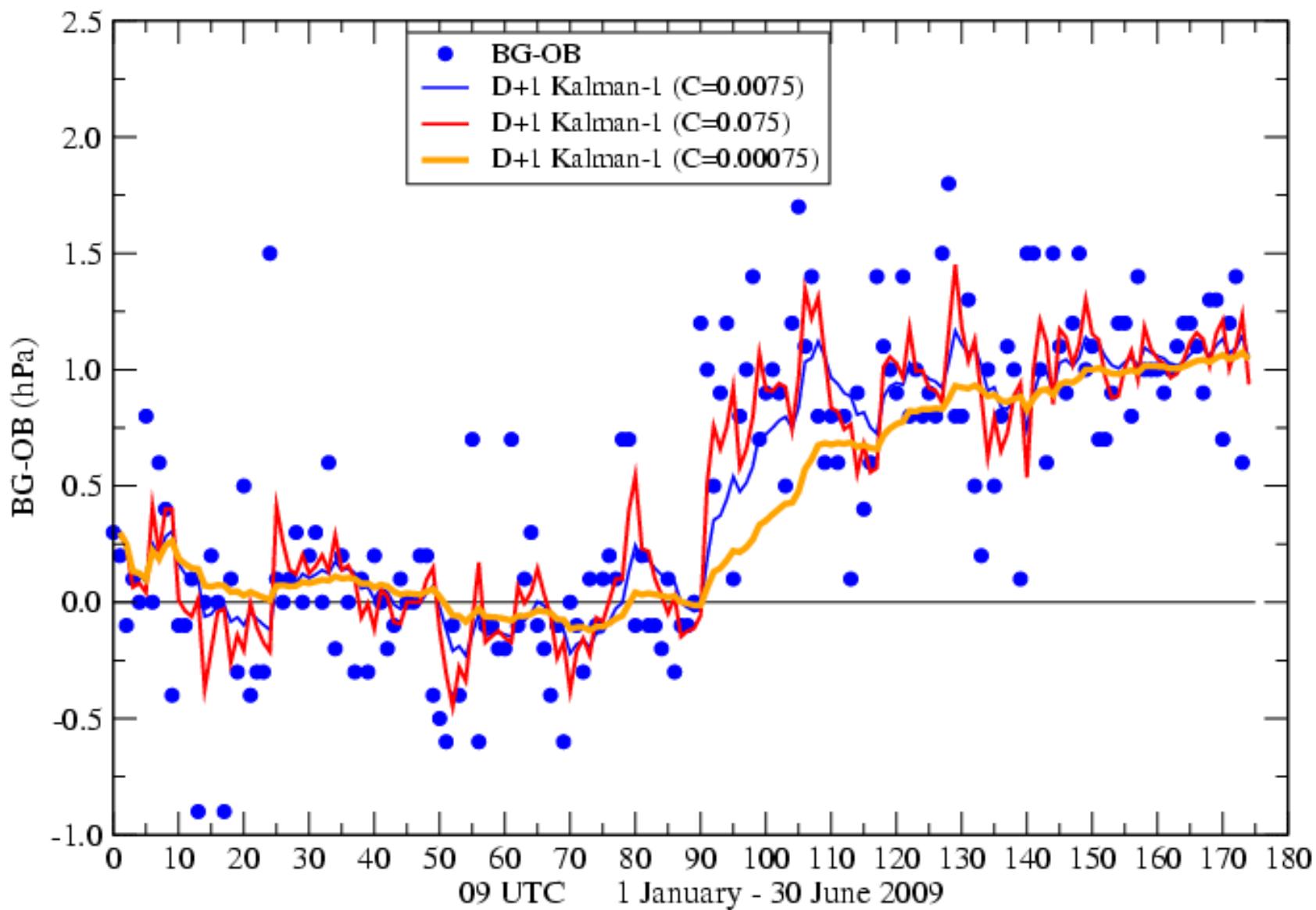
Kalman filtering of MSLP biases

SYNOP 03866, St Cathrine's Point, Isle of Wight ($C=0.0075$)



Kalman filtering of MSLP biases

SYNOP 03866, St Cathrine's Point, Isle of Wight (obs err=0.4)



Can we extract
probabilities, or at least
error bars, from the Kalman
filtered output?

The “Kalman gain”

$$k_t = \frac{P_{t/t-1} f_t}{r_t + f_t^T P_{t/t-1} f_t}$$

..where the nominator is the sum of the observation and model uncertainties and the updated model uncertainty is:

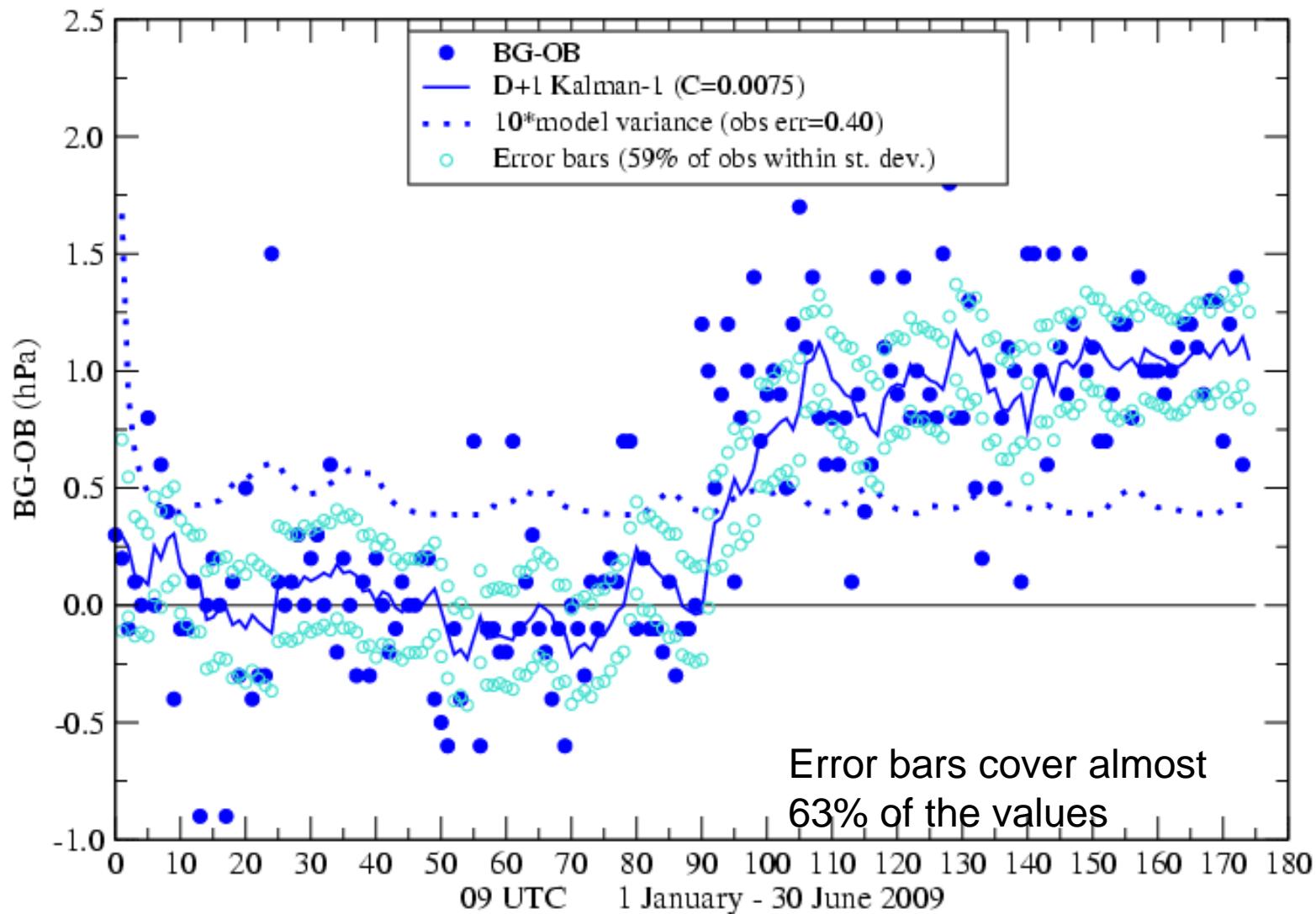
$$P_{t/t} = (I - k_t) P_{t/t-1} (I - k_t)^T + k_t r_t k_t^T$$

...which means that

$$(\text{total error})^2 = f_t^T P_{t/t} f_t$$

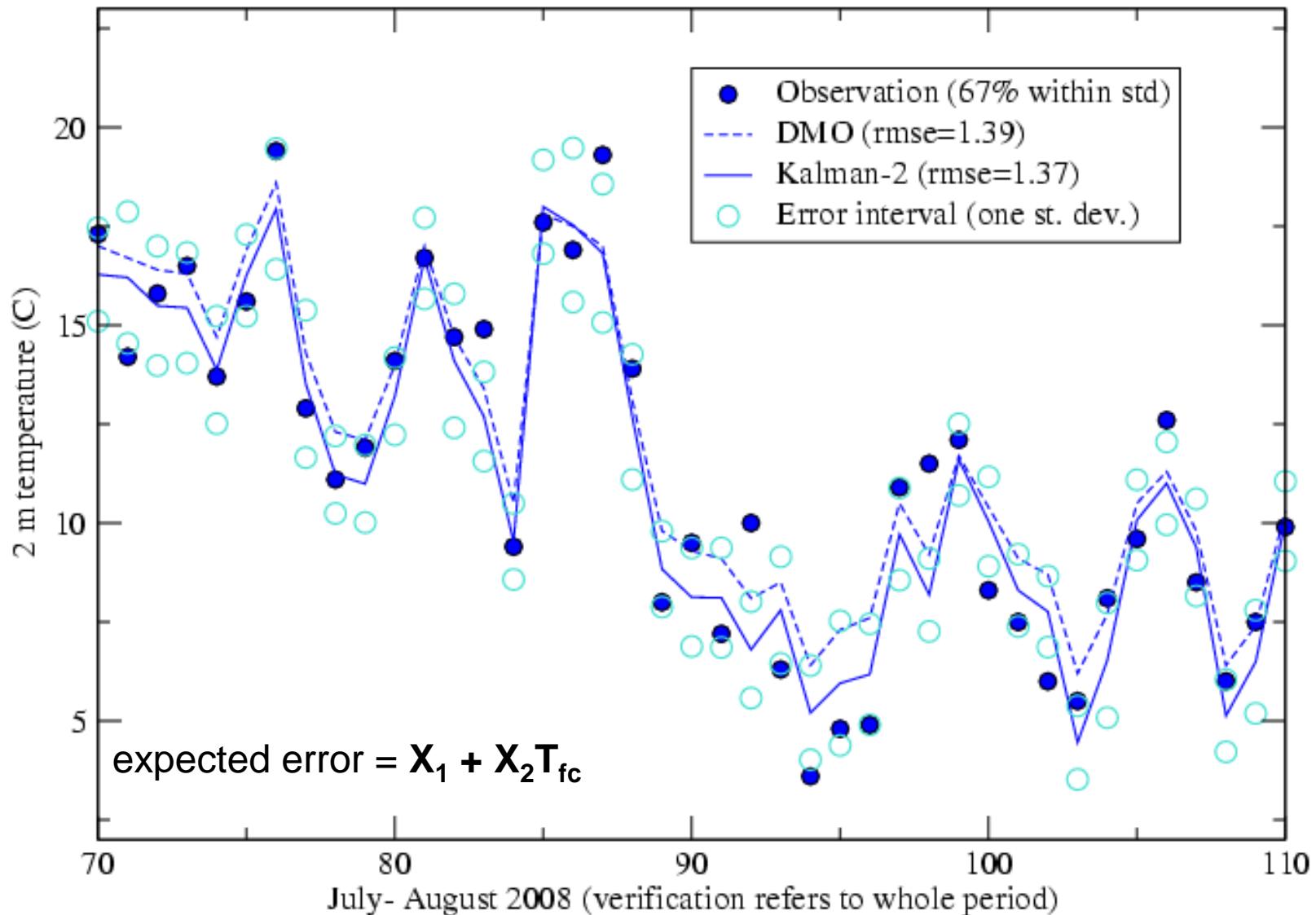
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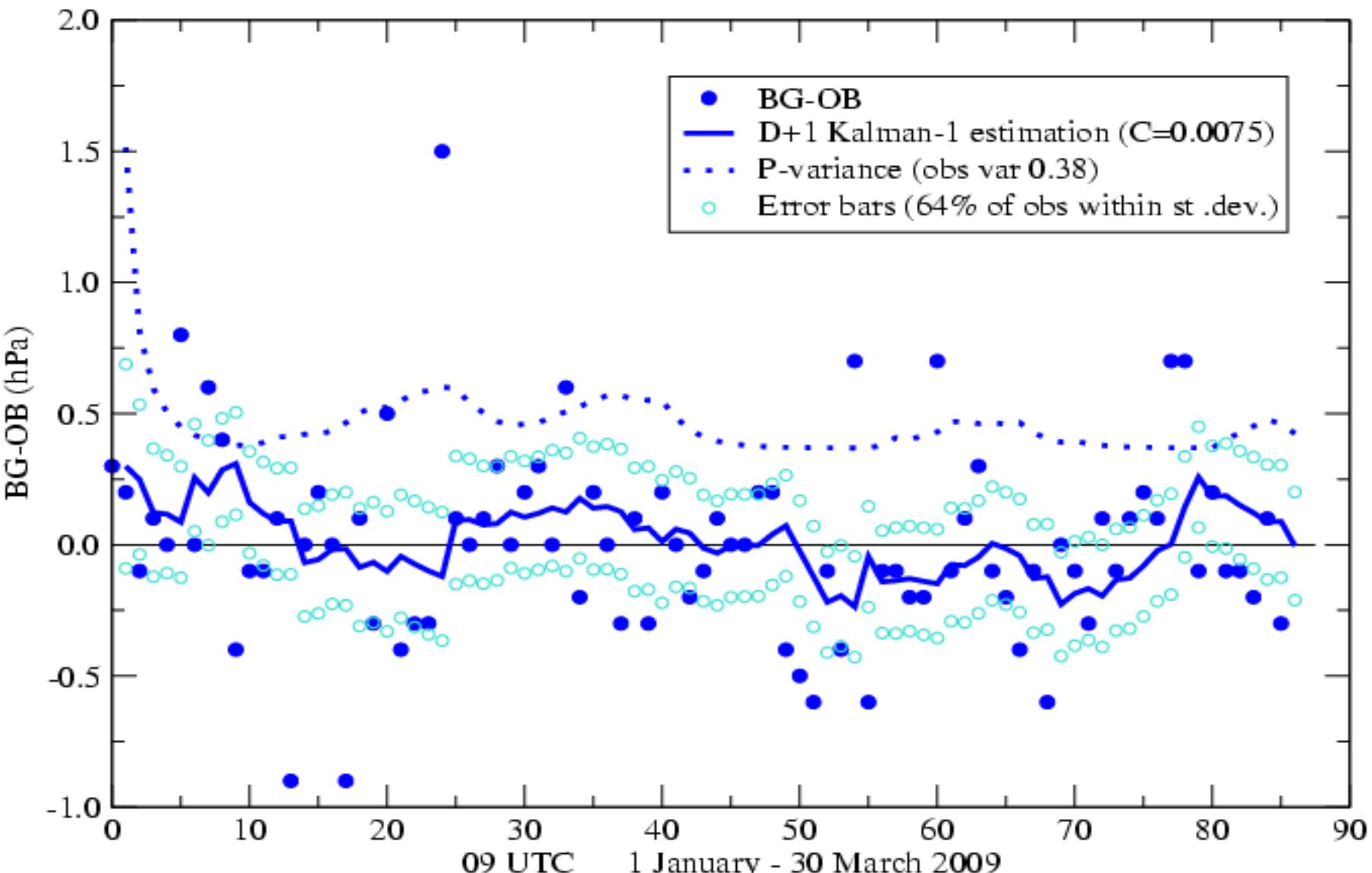
Kalman-2 filtered 2 m temperature forecasts

Volkel (NL) 12 UTC + 12 h ECMWF forecasts June 2008-April 2009



Kalman filtering of MSLP biases

SYNOP 03866, St Cathrine's Point, Isle of Wight



END